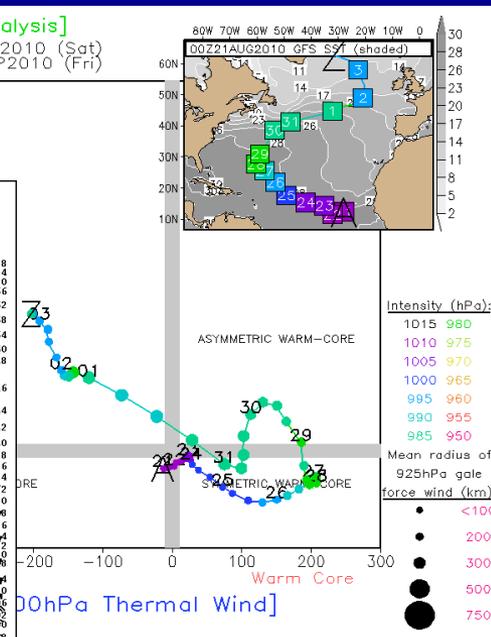
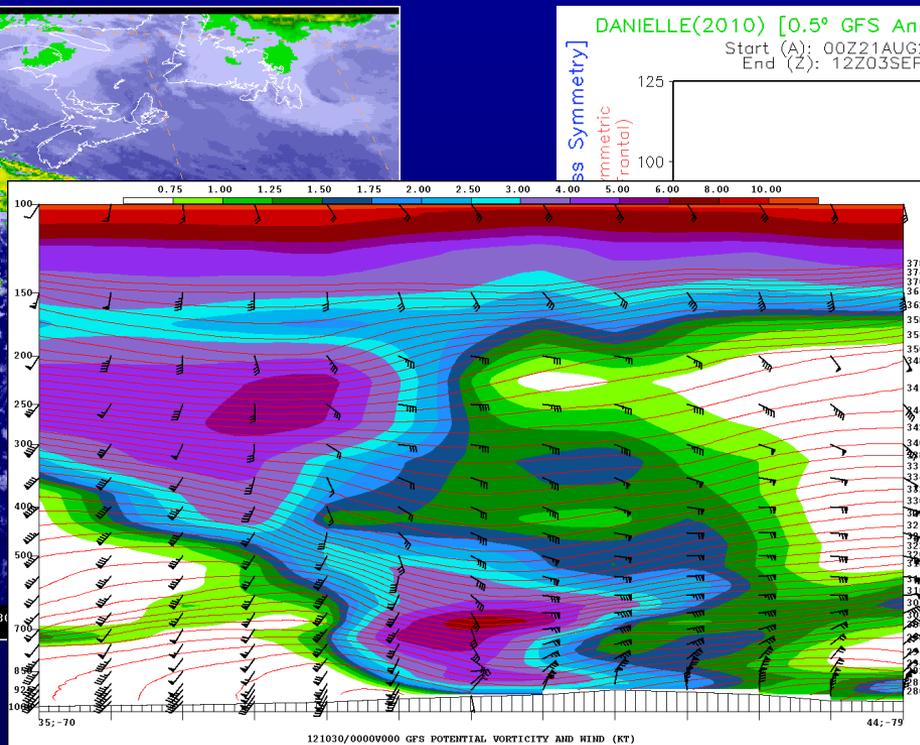
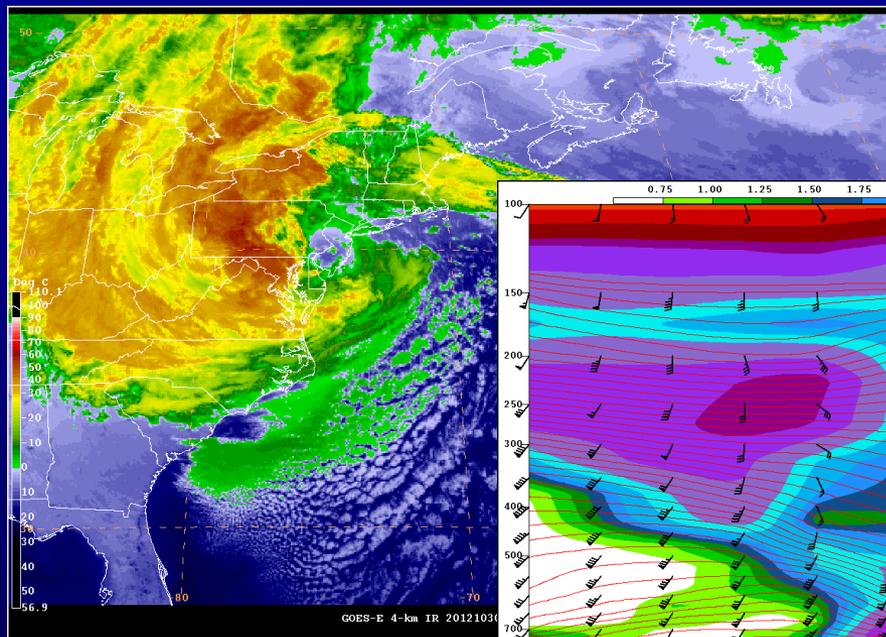




Extratropical Transition: Operational Challenges and Forecast Tools



Michael J. Brennan
National Hurricane Center
28 February 2018



Outline

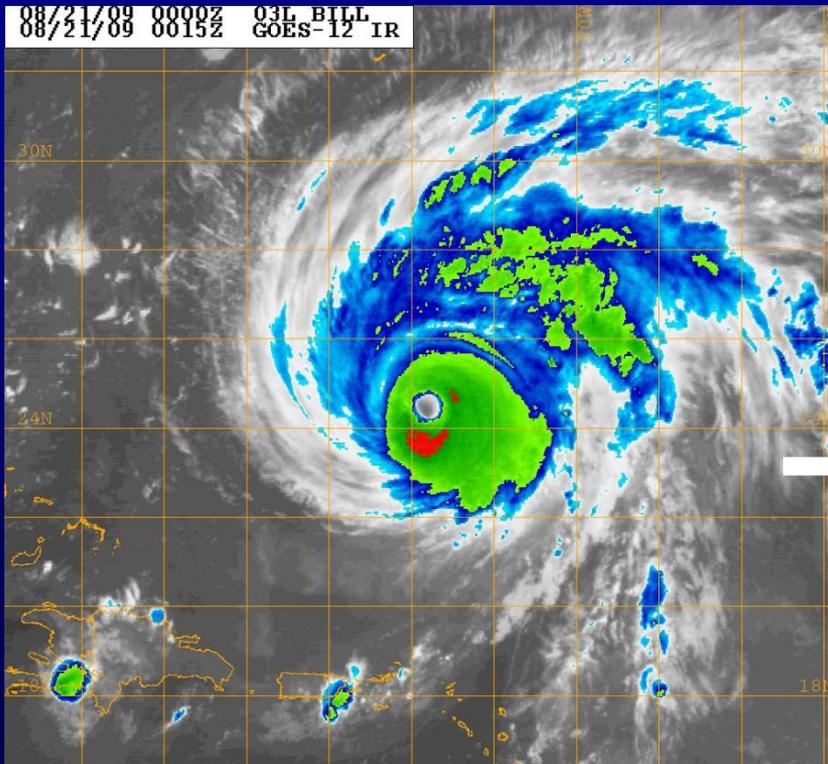


- Introduction
- Climatology
- Operational Procedures and Challenges
- Environmental/Structural Changes During ET
- Conceptual Models and Composites of ET
- Forecasting Tools
- Summary
- Exercise



Extratropical Transition (ET)

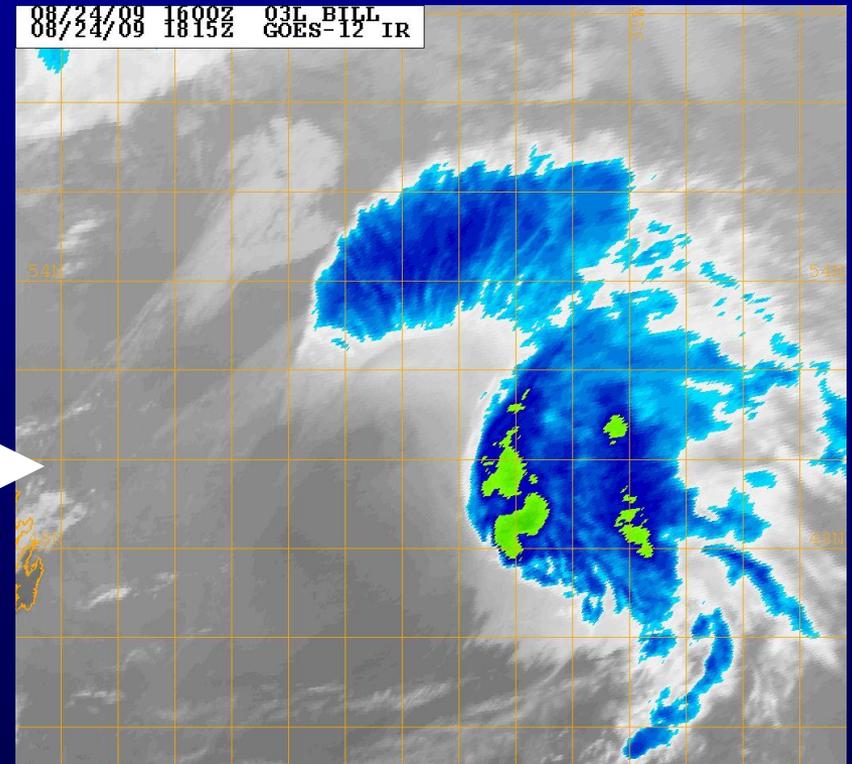
- “...a gradual process in which a tropical cyclone loses tropical characteristics and becomes more extratropical in nature.” (Jones et al. 2003)



Hurricane Bill 00 UTC 21 Aug 2009

Maximum Wind: 110 kt

34-kt wind radii: 225NE, 200SE, 120SW, 200NW



Extratropical Bill 18 UTC 24 Aug 2009

Maximum Wind: 60 kt

34-kt wind radii: 225NE, 275SE, 240SW, 150NW



Motivation – Why Study ET?



- Numerous tropical cyclones (TCs) undergo ET
 - Nearly half in some basins
- Structural changes during ET affect distribution and magnitude of hazardous weather
- Cyclones often affect land during or after ET
 - Northeast U.S., Atlantic Canada, Ireland, Great Britain, NW Europe
- Large impact on mid/high latitude shipping lanes
- A challenge to forecast!
 - High-impact, high-visibility events
 - Not always handled well by NWP models
 - Impacts on downstream longwave pattern and hemispheric predictability implications



Climatology of ET

Emphasis on North Atlantic Basin



Climatology Question



- In what TC basin does ET occur least often?
 1. North Atlantic
 2. Eastern North Pacific
 3. Western North Pacific
 4. Northern Indian Ocean

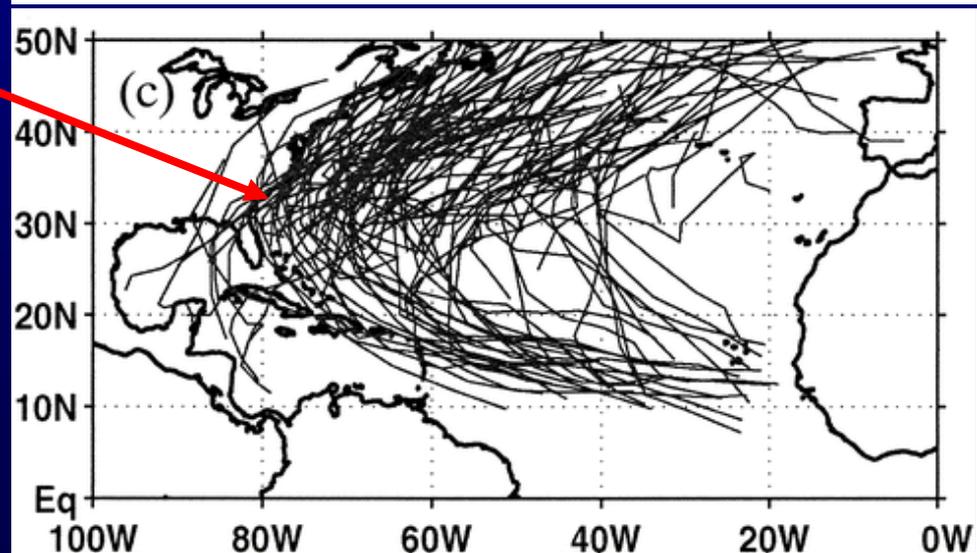
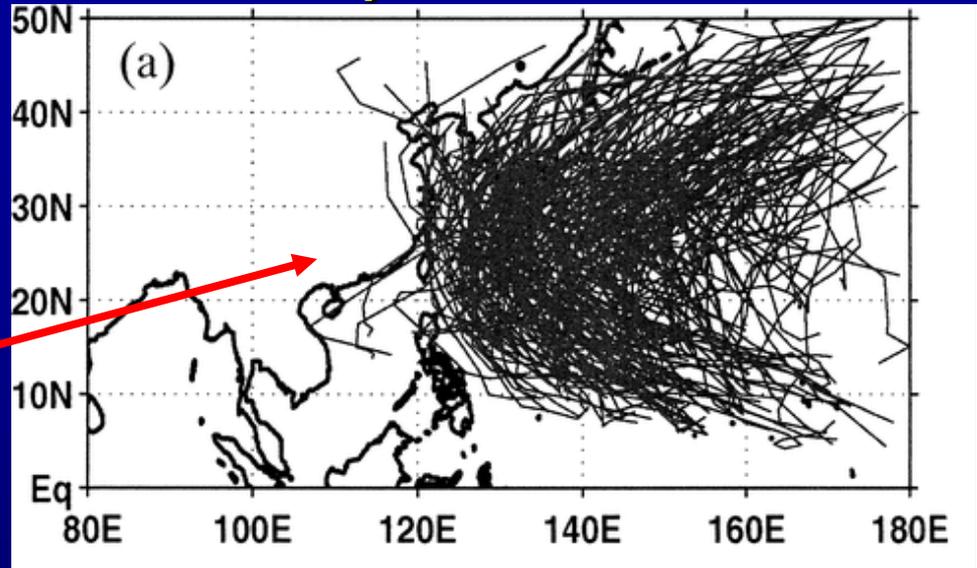


Climatology of ET

(Jones et al. 2003)



- ET occurs in almost every basin where TCs form
- Largest *number* of ET events occur in Northwest Pacific
- Largest *percentage* (~45%) of TCs undergo ET in North Atlantic
- Occurs least often in Northeast Pacific
 - Synoptic conditions and strong SST gradient not favorable for ET

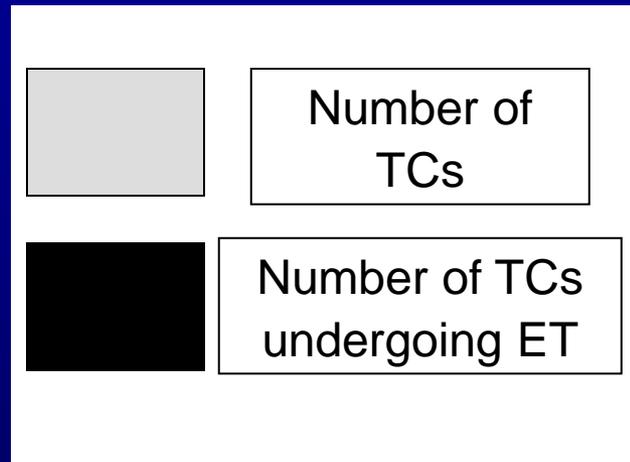
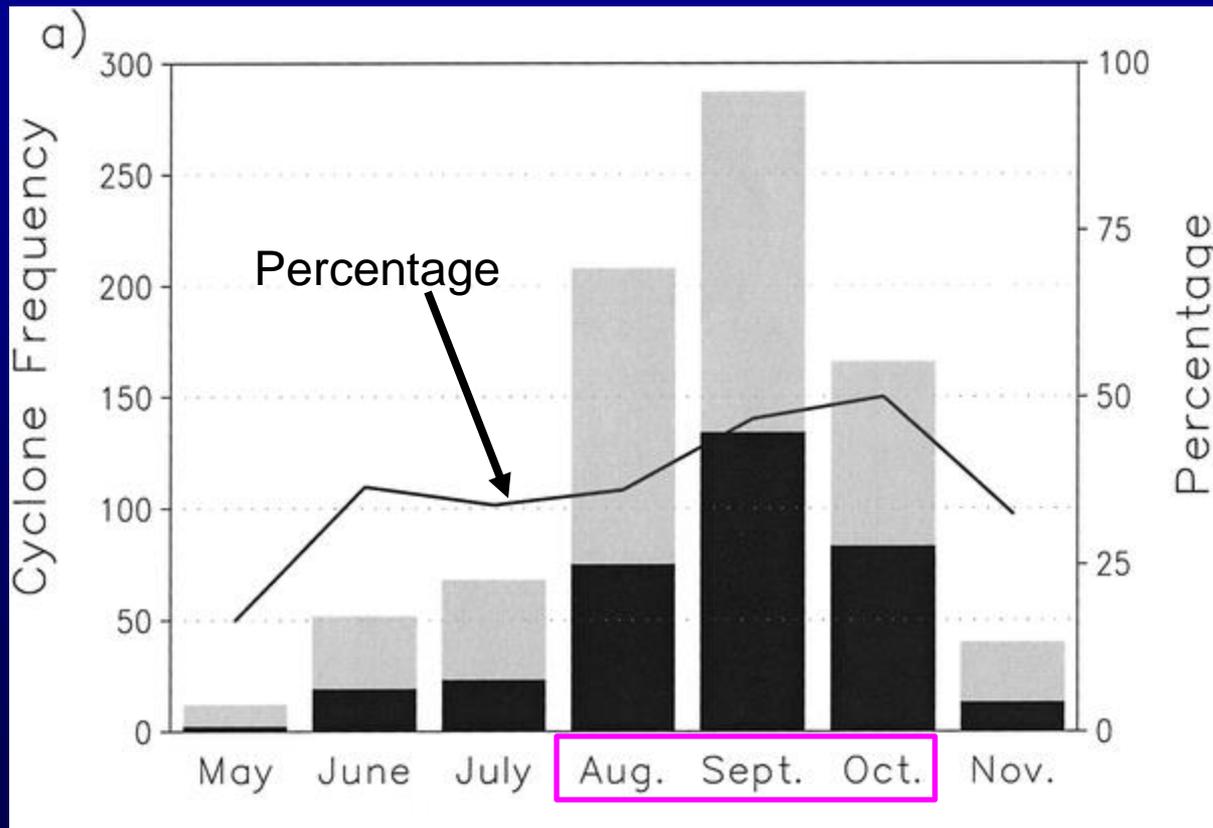




North Atlantic ET Climatology



Number and percentage of North Atlantic TCs undergoing ET by month 1899–1996

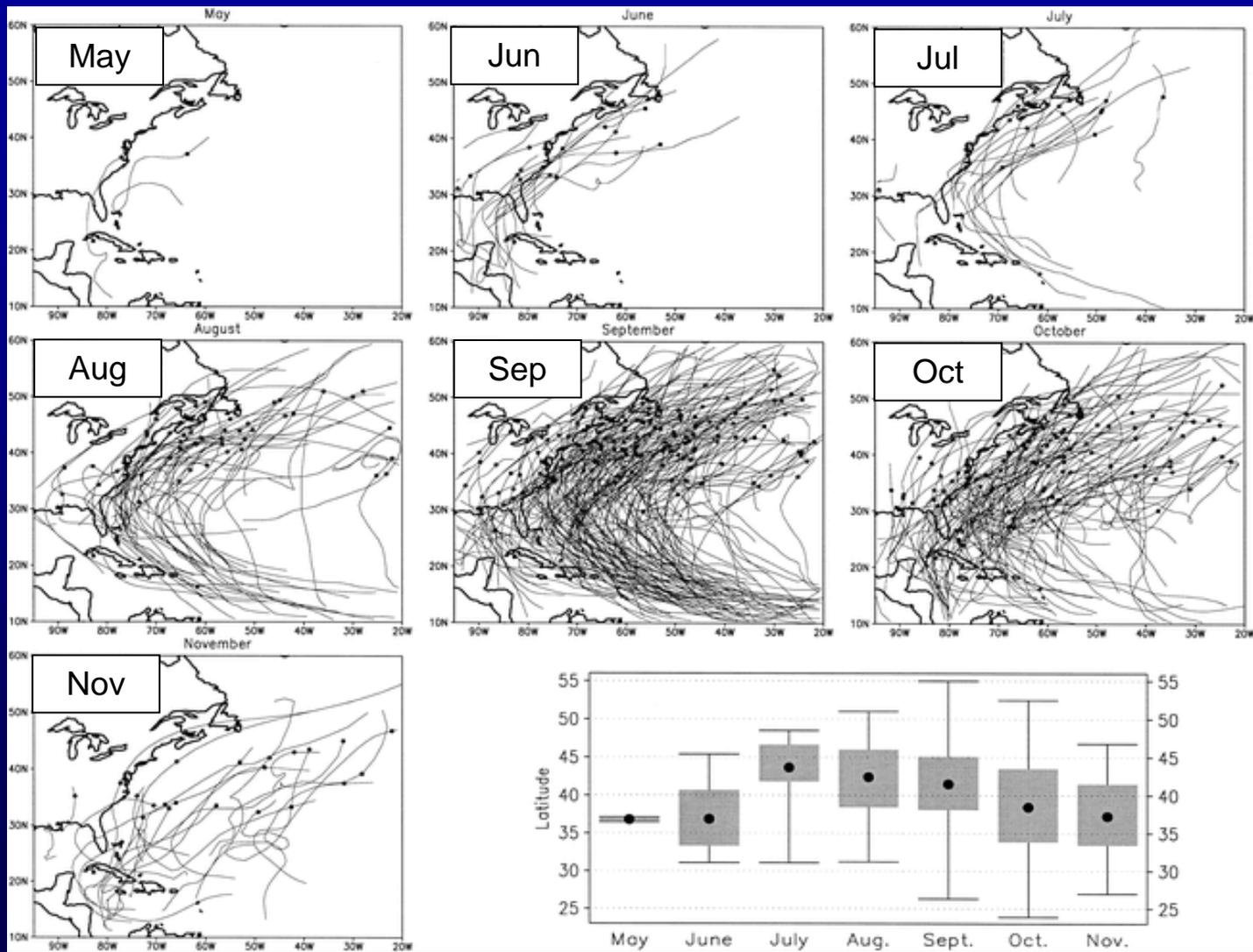


Largest number of Atlantic TCs undergo ET from August through October

Hart and Evans (2001) *Journal of Climate*



North Atlantic ET Climatology



Latitudinal distribution of ET in North Atlantic varies widely:

July through Sept: 40–50°N

Earlier and later in the season: 35–40°N



North Atlantic ET Climatology

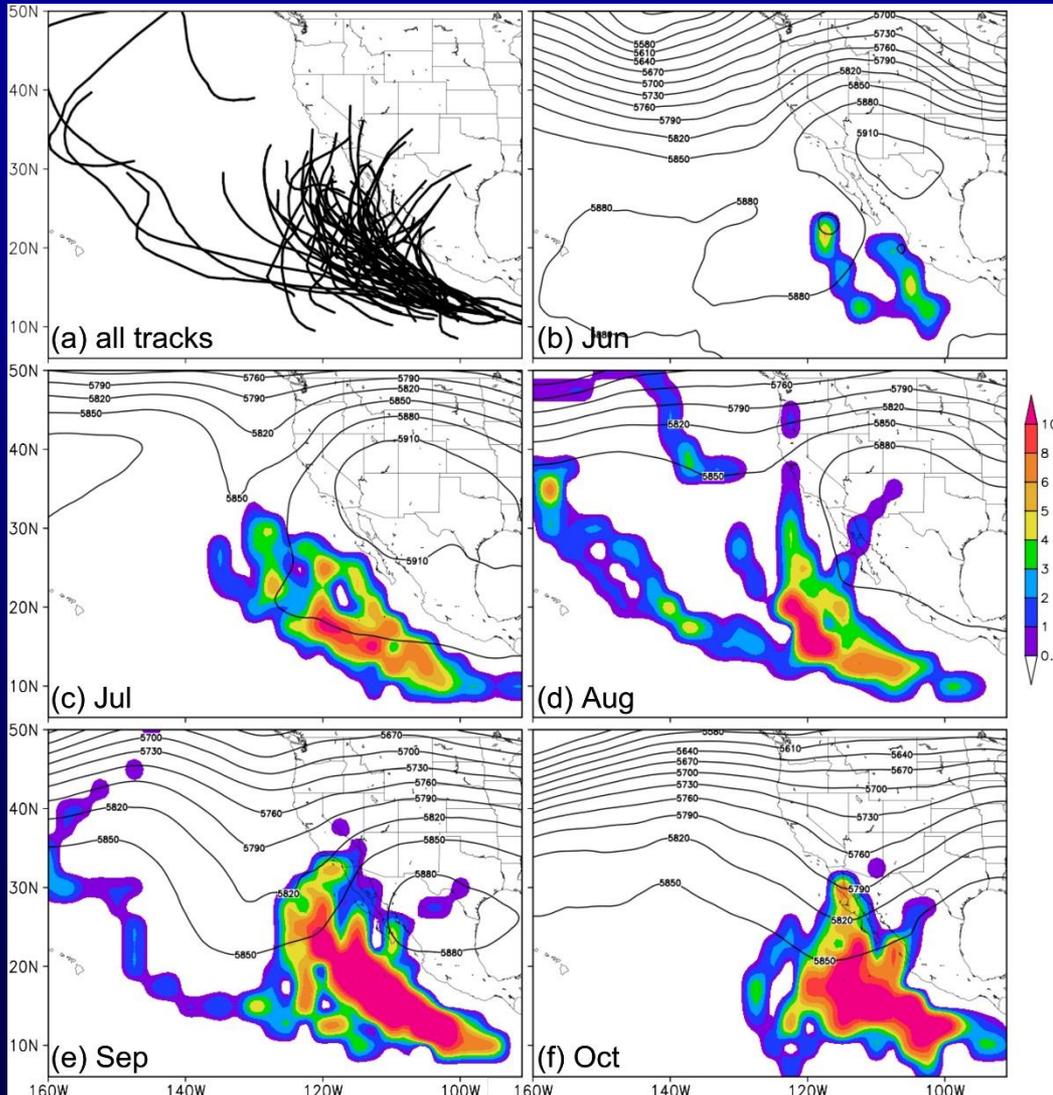
Hart and Evans (2001)



- 51% of storms undergoing ET from 1979–1993 intensified
 - Most initially formed in the deep Tropics (60% form south of 20°N)
- 42% of transitioning storms from 1979–1993 decayed during ET
 - Most form outside the deep Tropics (90% form north of 20°N)
- Initial deep tropical development more favorable for intensification later during ET
- TCs of more baroclinic origins more likely to weaken during ET



Eastern North Pacific ET Climatology (Wood and Ritchie 2014)



1971-2012 climatology based on phase space from JRA-55 reanalysis

About 9% of eastern North Pacific TCs underwent ET during this period (55 of 631)

Most ET event (65.5%) occur in Sep-Oct



Operational Procedures and Challenges



Operational Definitions



From NWS Directive 10-604

<http://www.weather.gov/directives/sym/pd01006004curr.pdf>

- **Tropical cyclone**

- Warm-core, non-frontal, synoptic-scale cyclone, originating over tropical or subtropical waters, with organized deep convection and a closed surface wind circulation about a well-defined center
- Includes tropical depressions (< 34 kt), tropical storms (34-63 kt), and hurricanes (64 kt or greater)

- **Subtropical cyclone**

- Non-frontal low pressure system that has characteristics of both tropical and extratropical cyclones, with a closed surface wind circulation about a well-defined center.
- Have organized moderate to deep convection, but lack a central dense overcast
- Derive a significant proportion of their energy from baroclinic sources, and are generally cold-core in the upper troposphere and often associated with an upper-level low or trough
- These systems generally have a radius of maximum winds relatively far from the center (usually greater than 60 n mi) compared to tropical cyclones, and generally have a less symmetric wind field and distribution of convection
- Includes subtropical depressions (< 34 kt) and subtropical storms (34-63 kt)



Operational Definitions

From NWS Manual 10-604

<http://www.weather.gov/directives/sym/pd01006004curr.pdf>



- **Post-Tropical cyclone**

- Generic term that describes a cyclone that no longer possesses sufficient characteristics to be considered a tropical cyclone
- Former tropical cyclones that have become fully extratropical and remnant lows are two specific classes of post-tropical cyclones

- **Remnant low**

- Post-tropical cyclone that no longer possesses the convective organization required of a tropical cyclone and has maximum sustained winds < 34 kt
- Most commonly applied to the nearly deep-convection-free swirls of stratocumulus in the eastern North Pacific

- **Extratropical cyclone**

- A cyclone of any intensity for which the primary energy source is baroclinic, i.e., results from the temperature contrast between warm and cold air masses



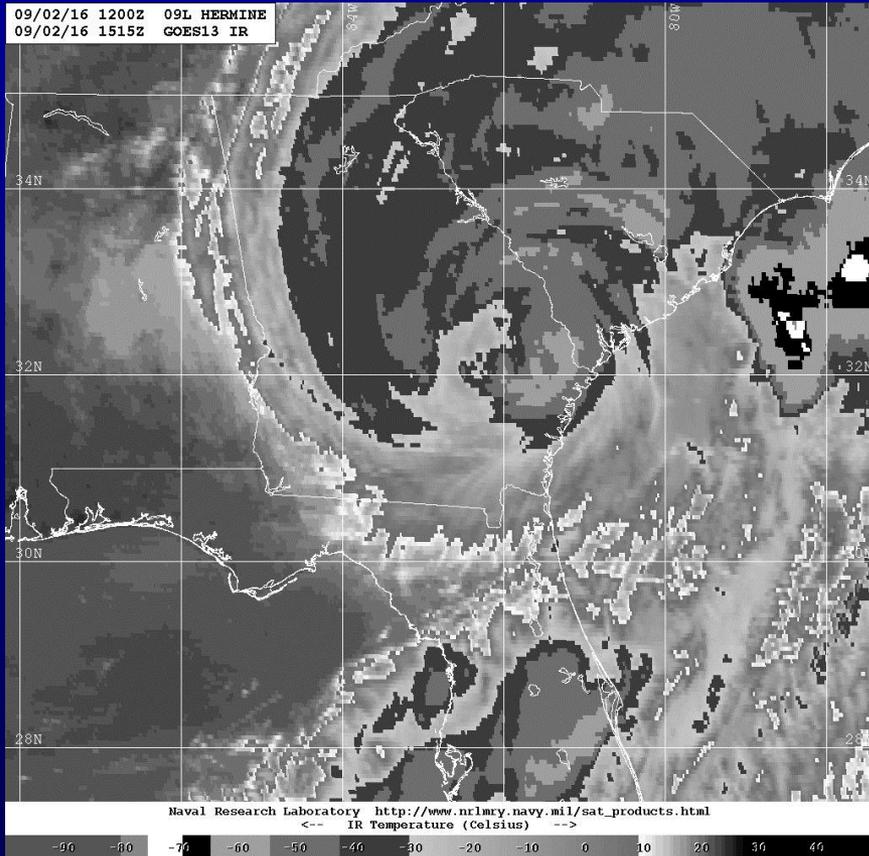
NHC Operational Procedures During ET



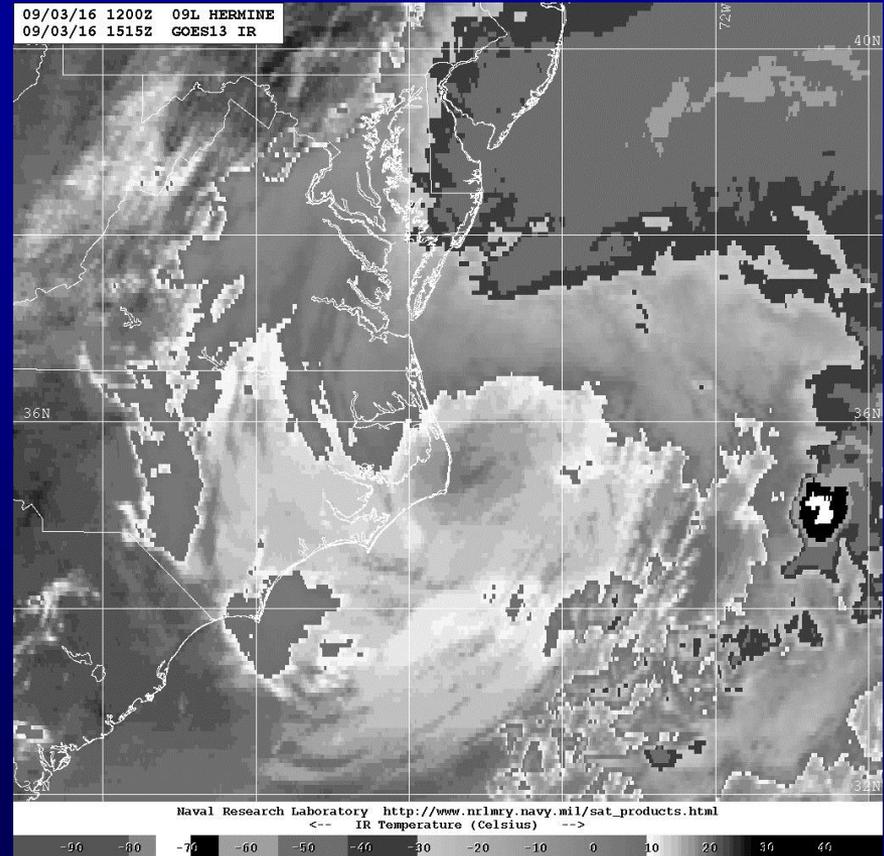
- Forecast periods designated **POST-TROPICAL** indicate the cyclone will no longer meet the definition of a tropical cyclone, but is not extratropical
- Forecast periods designated **POST-TROP/EXTRATROP** mean that the cyclone is expected to be extratropical
- Forecast track, intensity, and wind radii account for expected structural changes
- Forecasts collaborated with NWS partners and other nations (e.g., Canada, Bermuda, Ireland, UK)
- Advisories contain forecast information out to five days, including post-transition period, unless circulation expected to dissipate or be absorbed
- NHC will continue issuing advisories on a Post-Tropical Cyclone if it posts a significant threat to life and property and the transfer of responsibility would result in discontinuity of service
- Final NHC advisory indicates where to obtain further information on the system and highlights potential impacts



Post-Tropical Hermine (2016)



Tropical Storm Hermine
1515 UTC 2 Sep 2016



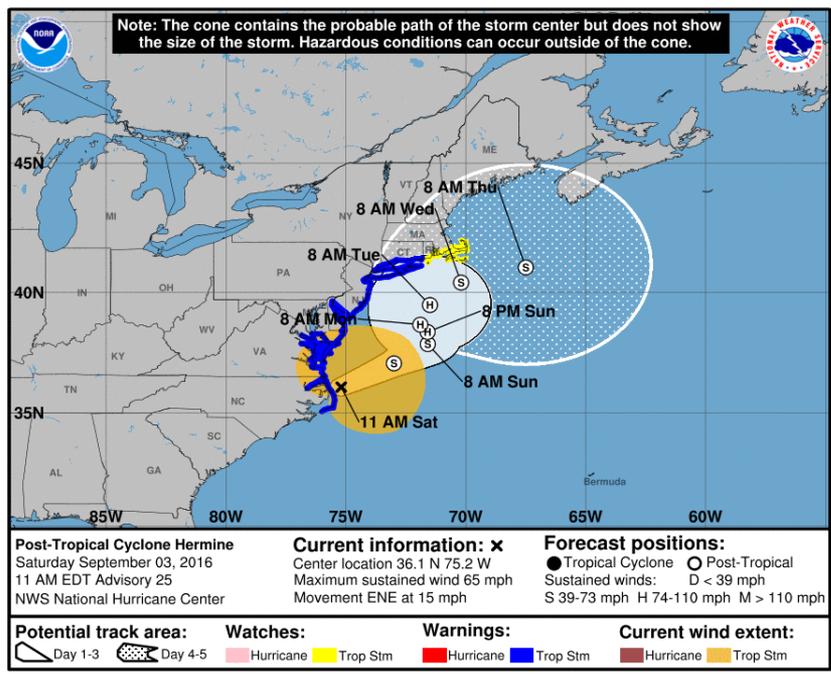
Post-Tropical Cyclone Hermine
1515 UTC 3 Sep 2016



Post-Tropical Hermine (2016)



Post-Tropical Cyclone Hermine Discussion Number 25
NWS National Hurricane Center Miami FL AL092016
1100 AM EDT Sat Sep 03 2016



Satellite imagery indicates that Hermine has become a **post-tropical cyclone**, with the coldest convective tops now located more than 200 n mi northeast of the exposed center. **Despite this change in structure, surface data from the Outer Banks indicate that some strong winds persist near the center, and the initial intensity is set to 55 kt for this advisory.** During the next 48 to 72 hours, Hermine will interact with a strong mid-latitude shortwave trough and **all of the global models show the system re-intensifying during that time and a redevelopment of a stronger inner core, albeit one situated underneath an upper-level low.** Regardless of its final structure, Hermine is expected to remain a dangerous cyclone through the 5 day period.

FORECAST POSITIONS AND MAX WINDS

INIT	03/1500Z	36.1N	75.2W	55	KT	65	MPH...POST-TROPICAL
12H	04/0000Z	37.1N	73.0W	60	KT	70	MPH...POST-TROPICAL
24H	04/1200Z	37.9N	71.6W	60	KT	70	MPH...POST-TROPICAL
36H	05/0000Z	38.4N	71.6W	65	KT	75	MPH...POST-TROPICAL
48H	05/1200Z	38.7N	71.9W	65	KT	75	MPH...POST-TROPICAL
72H	06/1200Z	39.5N	71.5W	65	KT	75	MPH...POST-TROPICAL
96H	07/1200Z	40.4N	70.2W	60	KT	70	MPH...POST-TROPICAL
120H	08/1200Z	41.0N	67.5W	50	KT	60	MPH...POST-TROPICAL



Operational Challenges

Is it Tropical or Extratropical?



- Cyclones don't always fit into discrete bins
 - Continuous spectrum exists between cyclones driven largely by diabatic heating (TCs) and those driven by baroclinic processes (extratropical cyclones)
 - Cyclones undergoing ET may derive energy from both diabatic and baroclinic processes
 - This often makes the decision of when to declare a TC “Post-Tropical” or “Extratropical” quite subjective
- When should NHC stop writing advisories on a tropical cyclone undergoing ET?
 - Decision has large impact on public perception of the threat associated with a system
 - Impacts watch/warning process
 - Requires large amounts of coordination for land and marine forecasts

If you get hit by 60 kt winds, 6 ft above normal tides, and/or 12 in of rain, does the exact nature of the system really matter?



Operational Challenges



- Will cyclone undergo ET?
- Will ET be continuous or be interrupted?
- Will cyclone intensify or weaken during ET?
- Will it re-intensify after transition?
- Will the TC become a discrete extratropical cyclone or be absorbed by a larger extratropical cyclone or front?



Question



- Which of the following structural changes does *not* typically occur during ET?
 1. Increase in the size of 34-kt wind field
 2. Redistribution of rainfall well away from the center
 3. Increase in maximum winds
 4. Decrease in deep convection
 5. Development of frontal boundaries



Operational Challenges

Sensible Weather Impacts



- Structural and intensity changes can occur rapidly during ET
- Maximum wind often decreases, but not always
- Overall impact of high winds may increase dramatically due to expansion of wind field
 - Major implications for land (wind) and marine interests (wind, waves, swell) as area of hazardous weather can become very large
- Drastic redistribution of precipitation
 - Tropical moisture can be carried far from its origins and interaction with synoptic-scale forcing, jets, and fronts can result in large rainfall totals
- Cyclones undergoing ET often interact with continental landmasses
 - Often source of cool, dry, continental airmass
 - Interactions with terrain



Operational Challenges

Model Forecasts



- Model forecasts of ET can be highly variable due to
 - Differences in initial cyclone structure and intensity
 - Differences in evolution of upper-tropospheric flow and other features of interest
 - Forecast track of TC, interaction with SST gradients, other features
- These factors determine if and when a TC begins ET, completes ET, and changes in intensity, structure and track



Hurricane Sandy (2012)

00 UTC 5 October

Intensity 85 kt

Pressure 964 mb

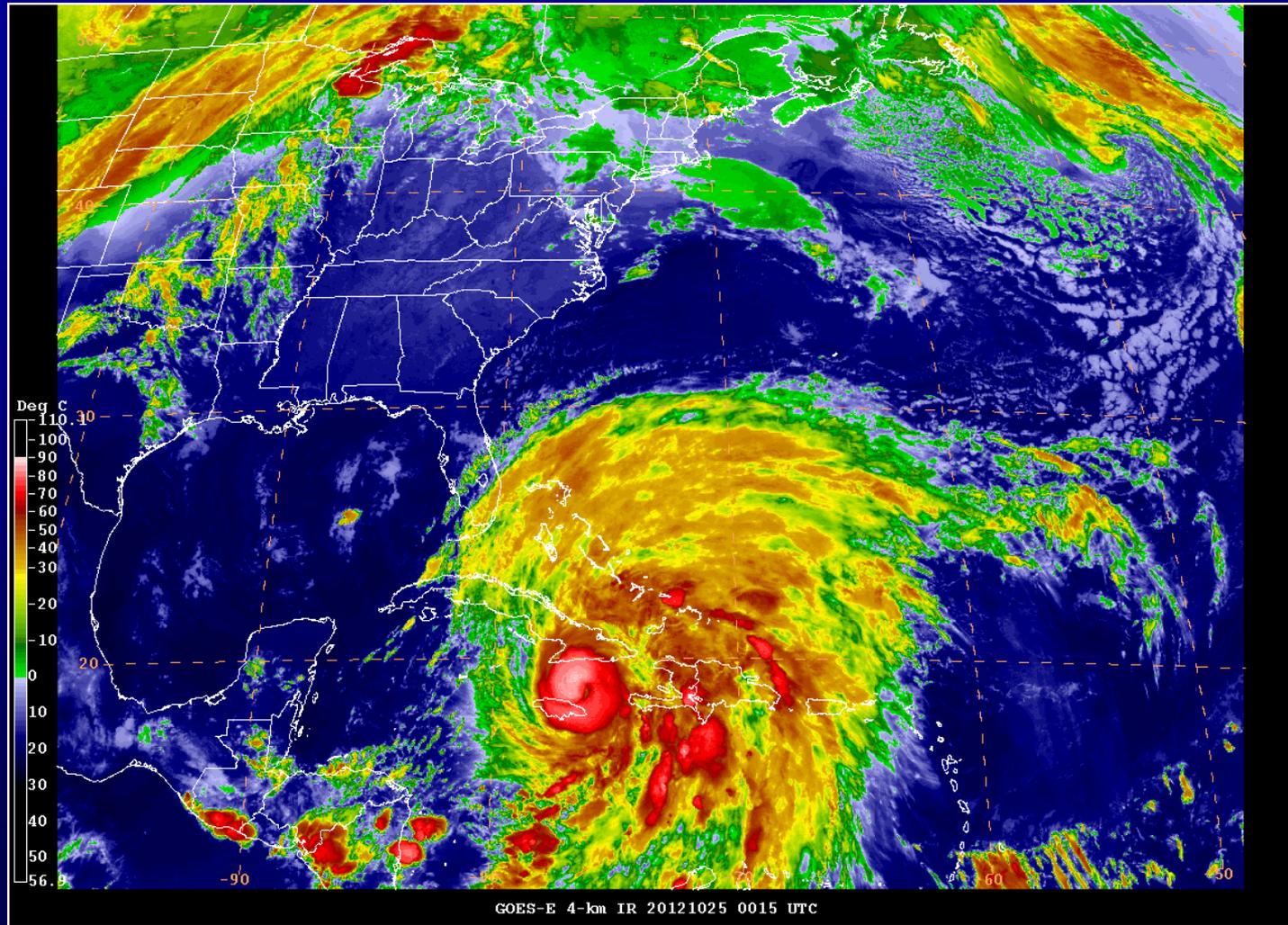
34 kt wind radii:

NE 180

SE 240

SW 70

NW 70





Hurricane Sandy (2012)

06 UTC 5 October

Intensity 100 kt

Pressure 954 mb

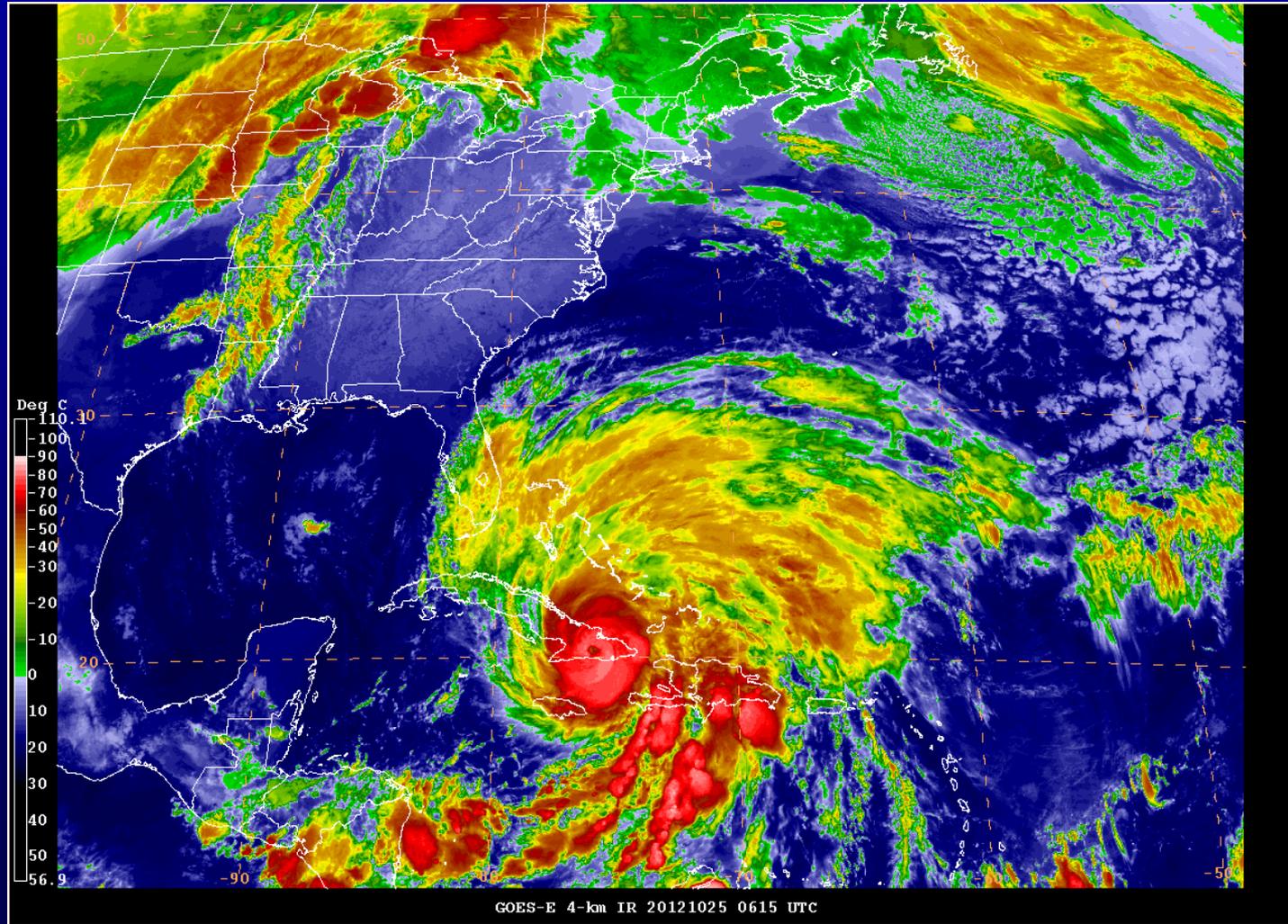
34 kt wind radii:

NE 240

SE 240

SW 70

NW 120





Hurricane Sandy (2012)

12 UTC 5 October

Intensity 95 kt

Pressure 966 mb

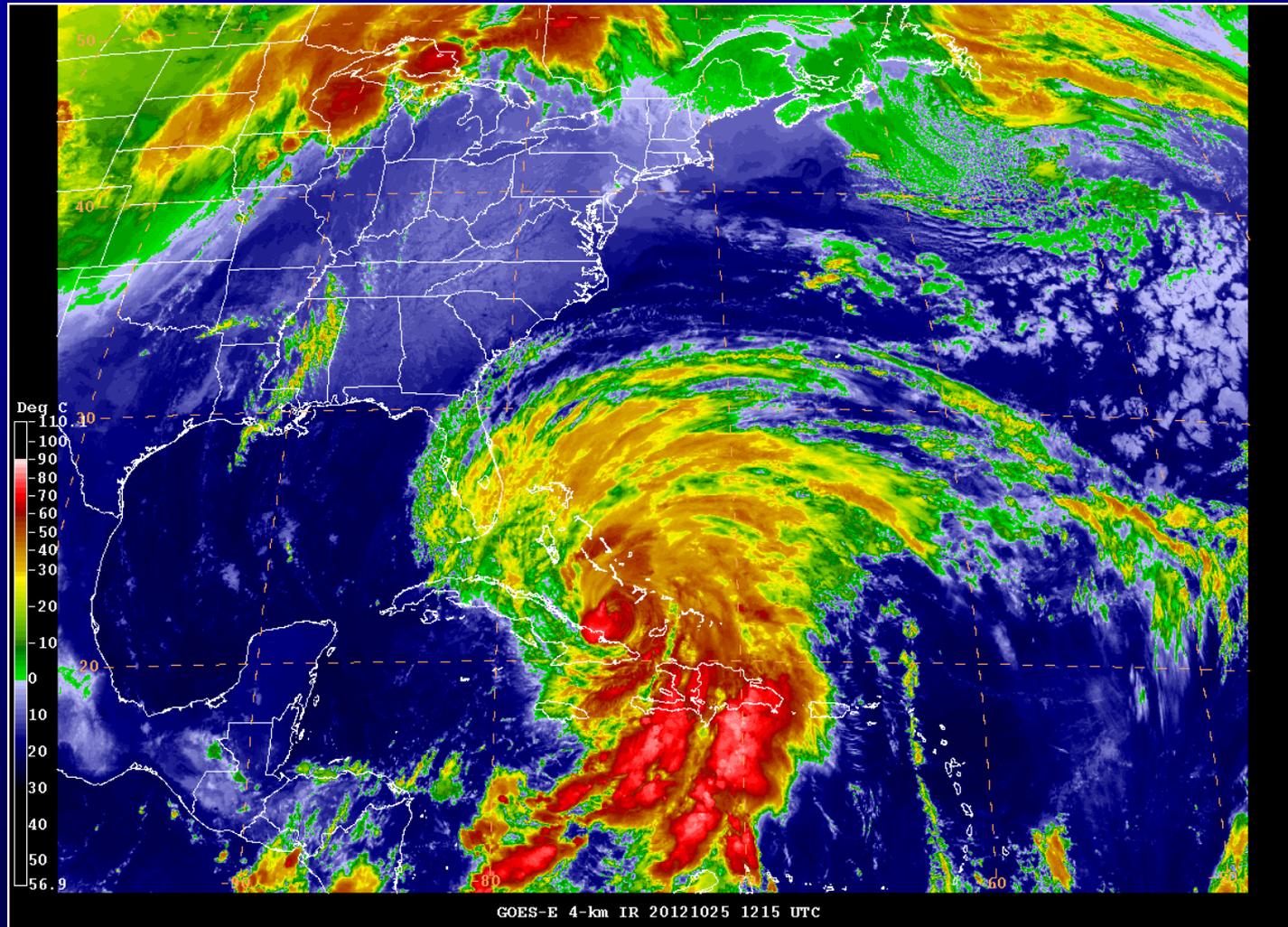
34 kt wind radii:

NE 240

SE 240

SW 70

NW 240





Hurricane Sandy (2012)

18 UTC 5 October

Intensity 90 kt

Pressure 963 mb

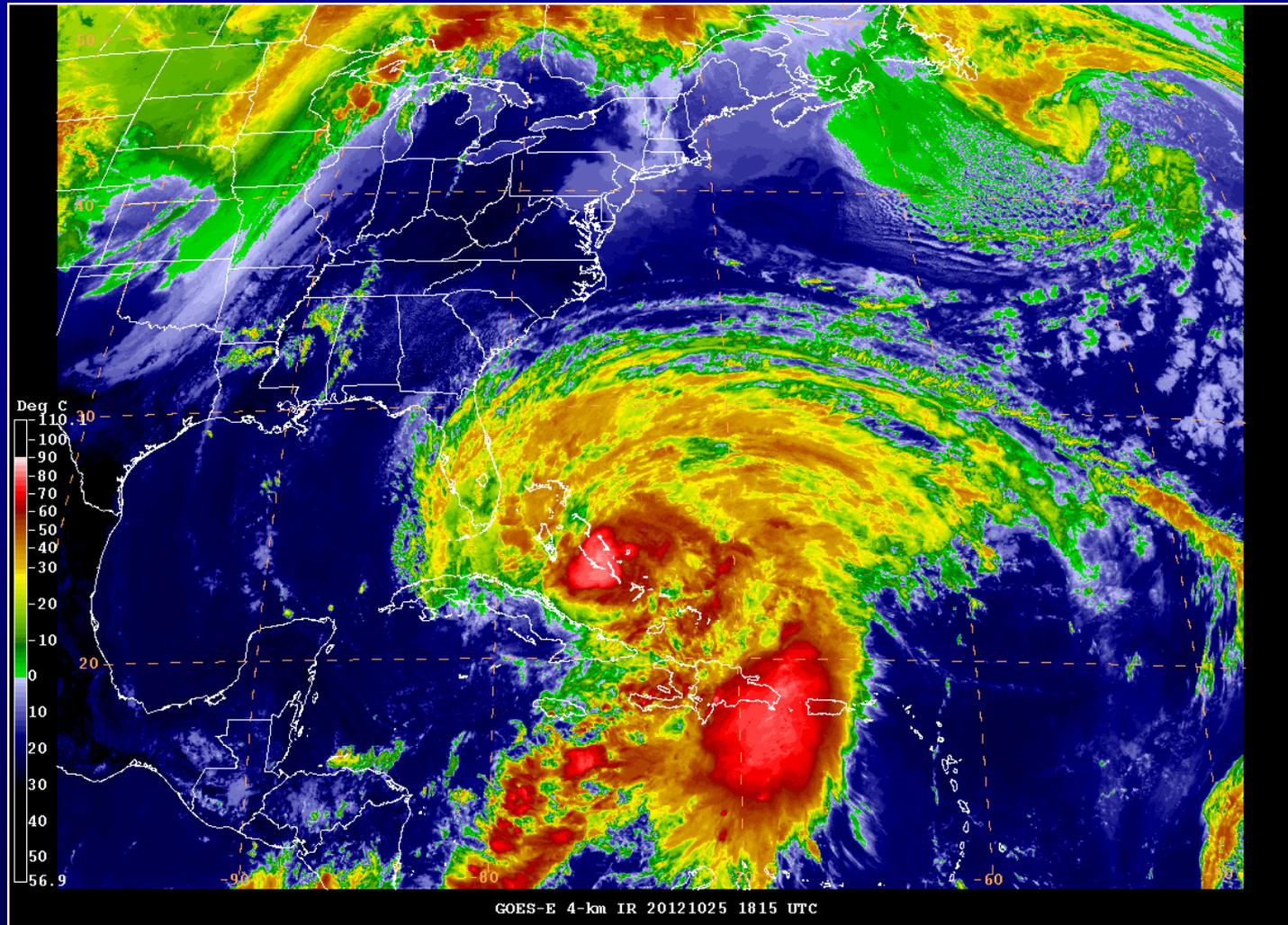
34 kt wind radii:

NE 270

SE 270

SW 120

NW 270





Hurricane Sandy (2012)

00 UTC 6 October

Intensity 75 kt

Pressure 965 mb

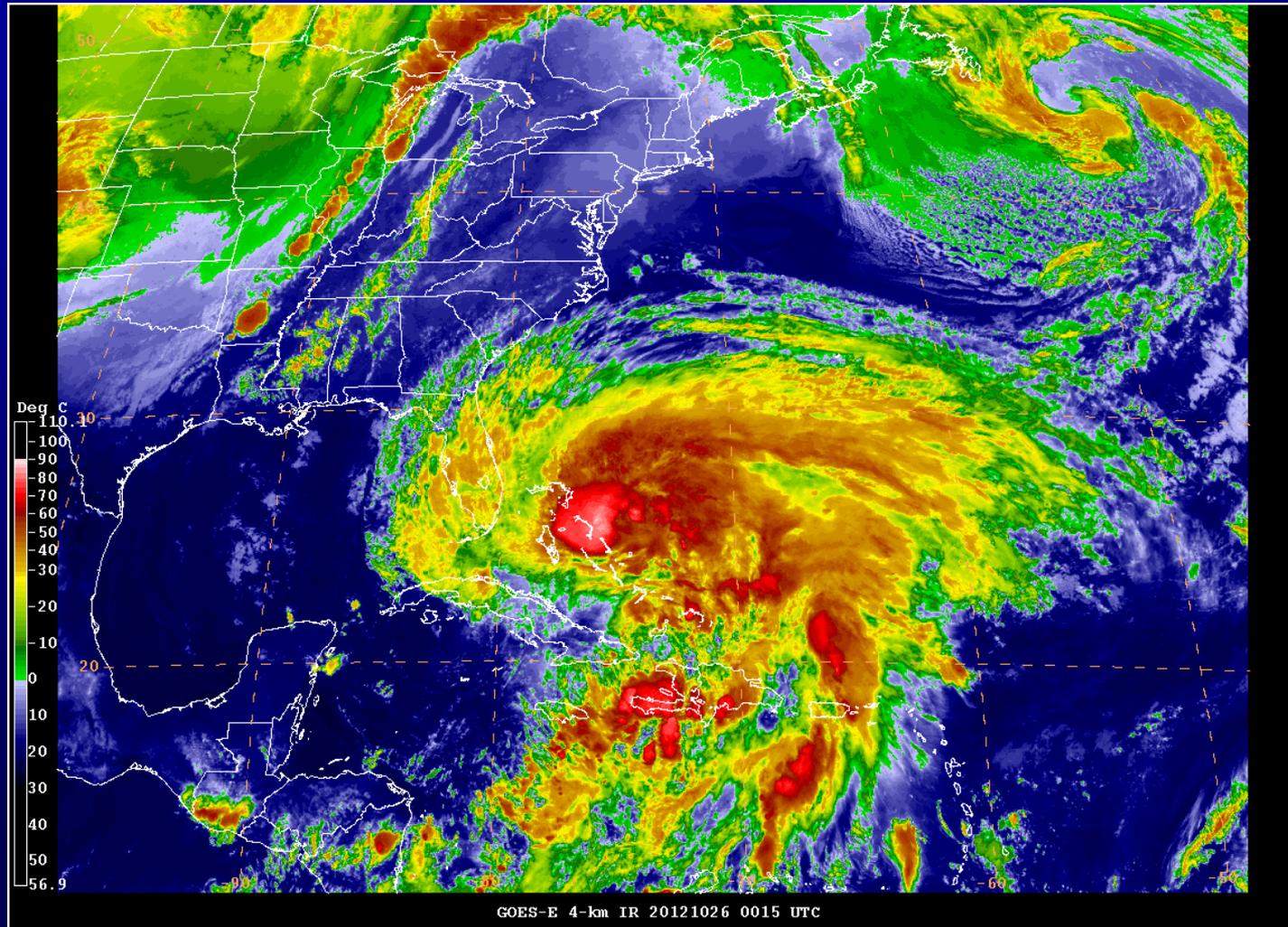
34 kt wind radii:

NE 300

SE 300

SW 160

NW 270





Hurricane Sandy (2012)

06 UTC 6 October

Intensity 70 kt

Pressure 968 mb

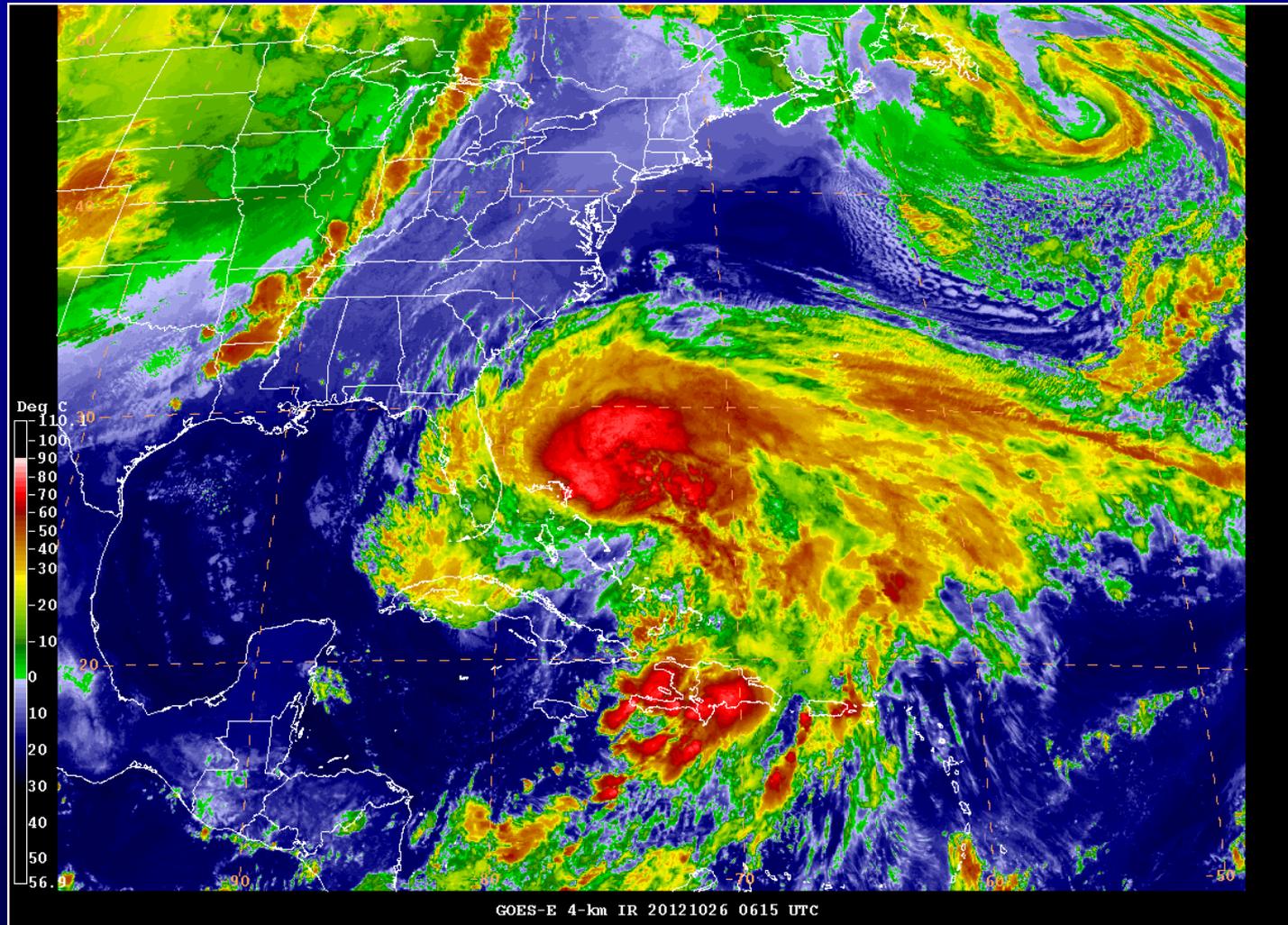
34 kt wind radii:

NE 300

SE 300

SW 160

NW 240





Hurricane Sandy (2012)

12 UTC 6 October

Intensity 65 kt

Pressure 970 mb

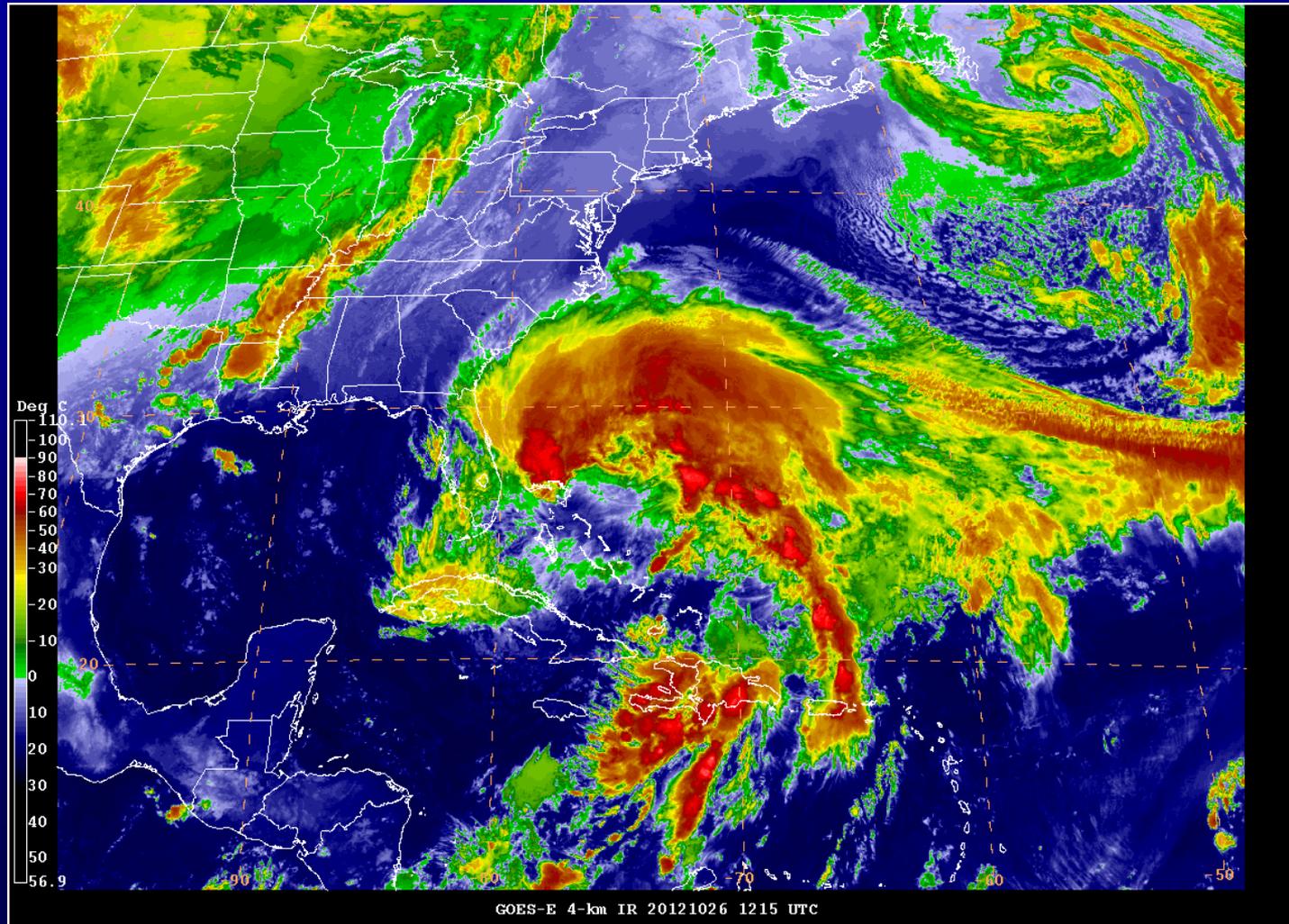
34 kt wind radii:

NE 360

SE 240

SW 170

NW 240





Hurricane Sandy (2012)

18 UTC 6 October

Intensity 65 kt

Pressure 971 mb

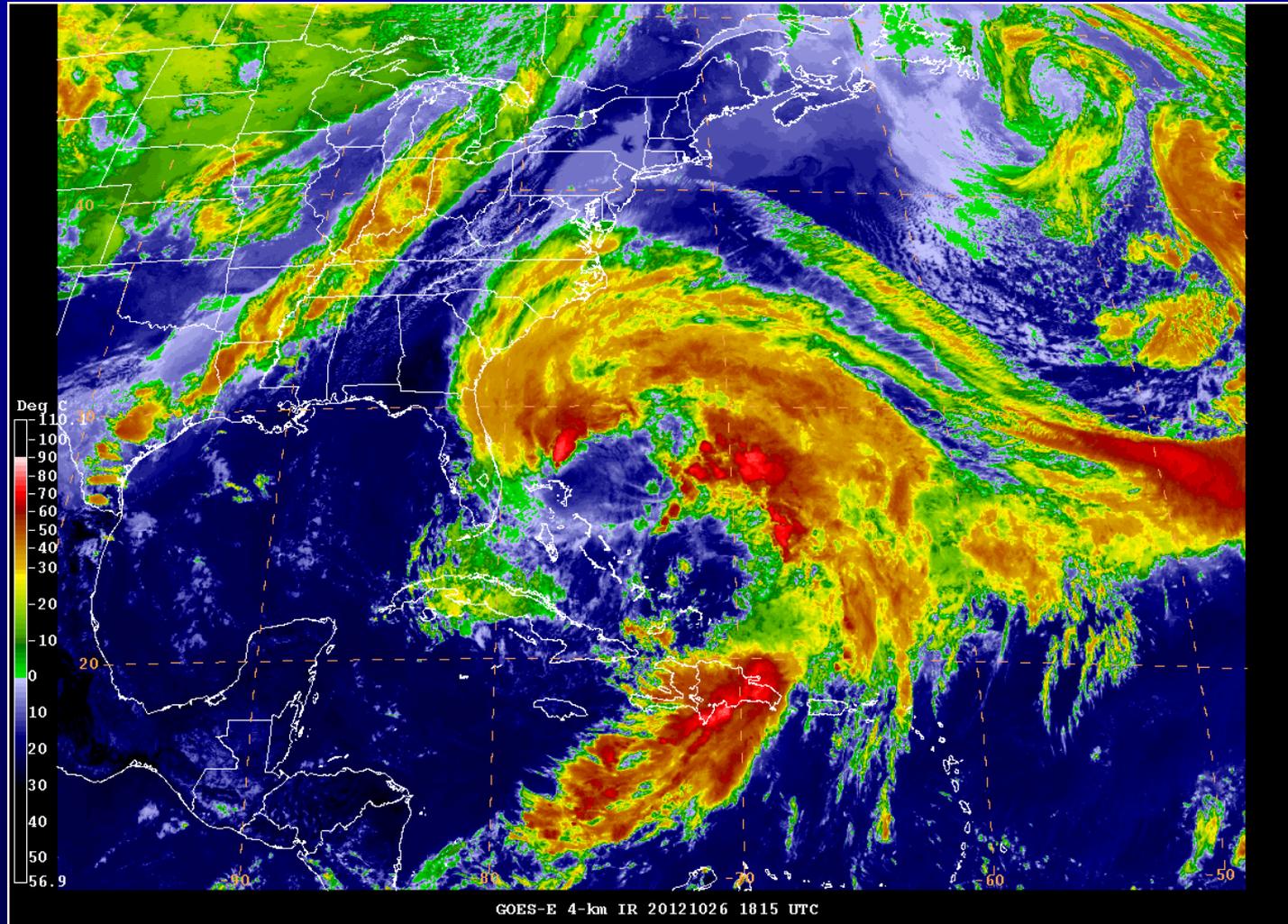
34 kt wind radii:

NE 400

SE 210

SW 170

NW 240





Hurricane Sandy (2012)

00 UTC 7 October

Intensity 60 kt

Pressure 969 mb

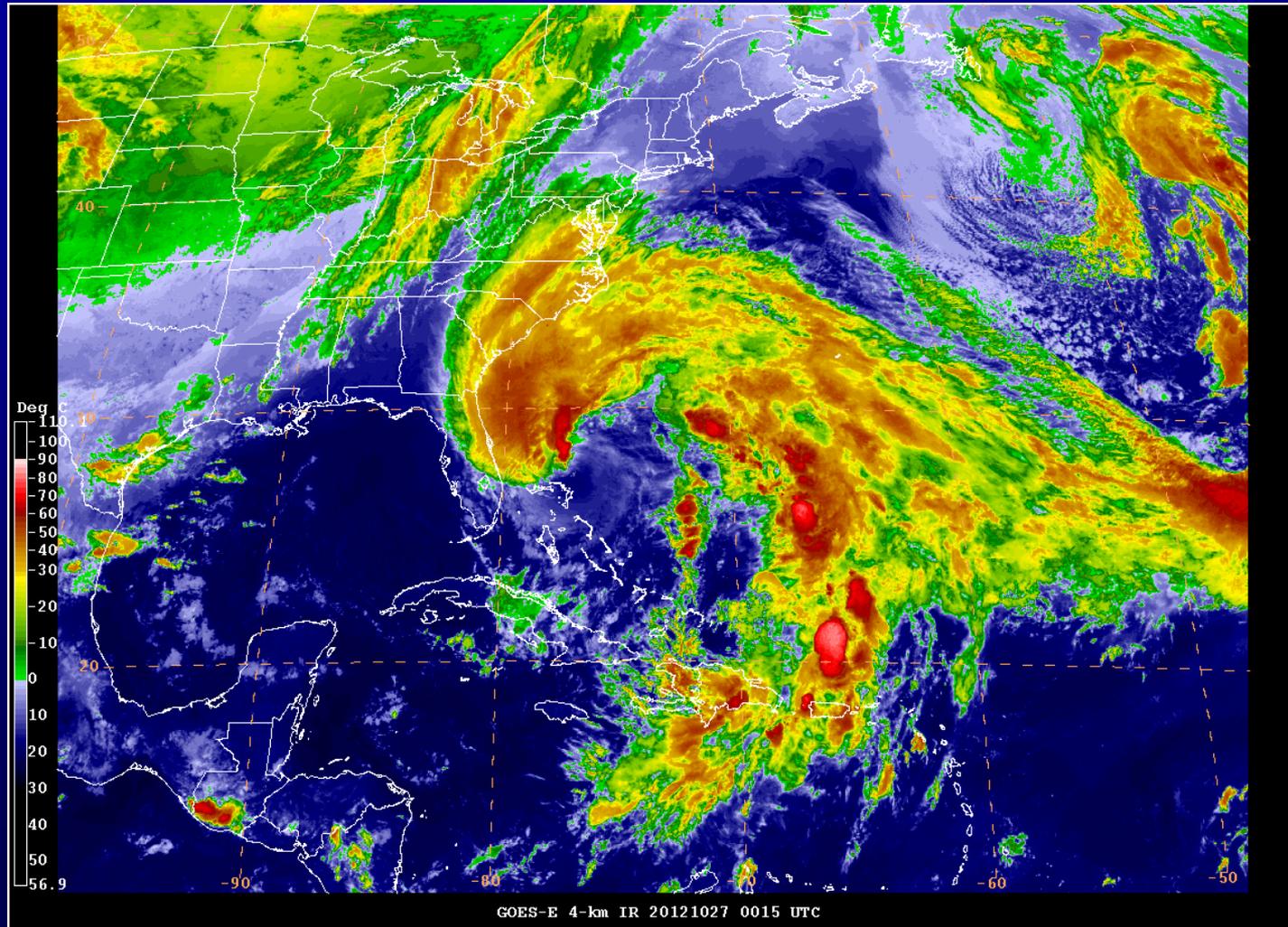
34 kt wind radii:

NE 450

SE 210

SW 180

NW 270





Hurricane Sandy (2012)

06 UTC 7 October

Intensity 60 kt

Pressure 968 mb

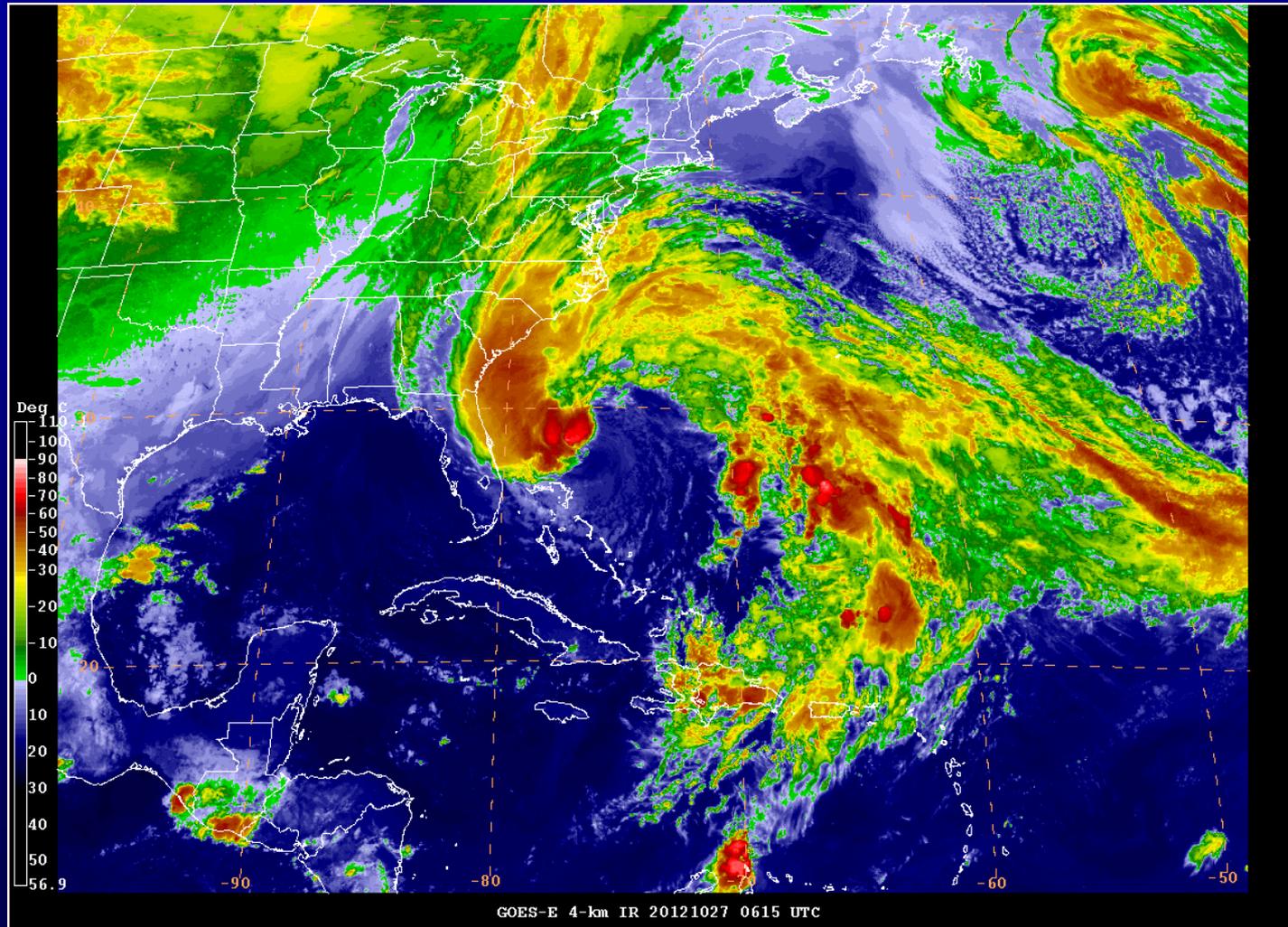
34 kt wind radii:

NE 450

SE 260

SW 180

NW 280





Hurricane Sandy (2012)

12 UTC 7 October

Intensity 70 kt

Pressure 956 mb

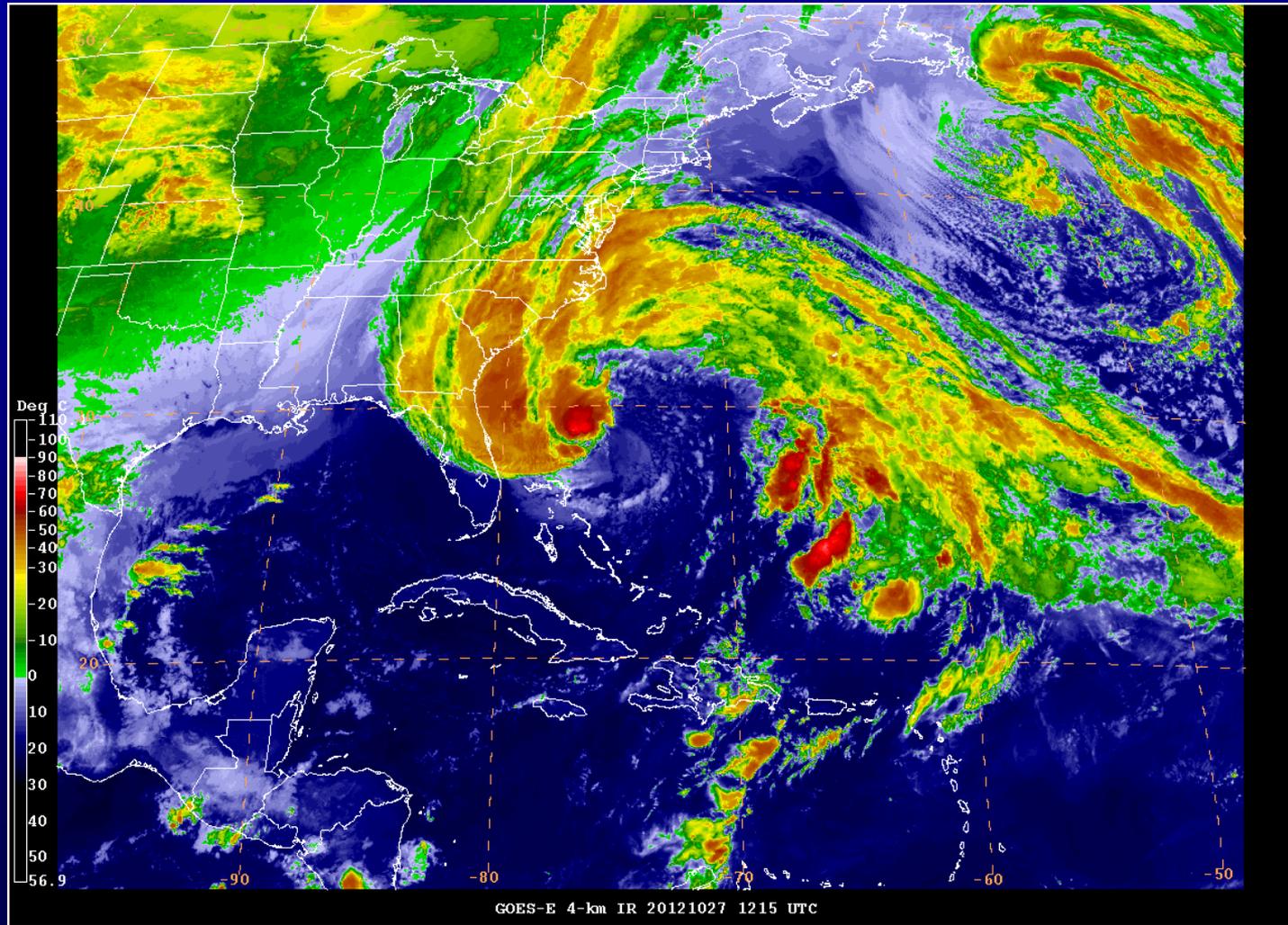
34 kt wind radii:

NE 450

SE 300

SW 210

NW 280





Hurricane Sandy (2012)

18 UTC 7 October

Intensity 70 kt

Pressure 960 mb

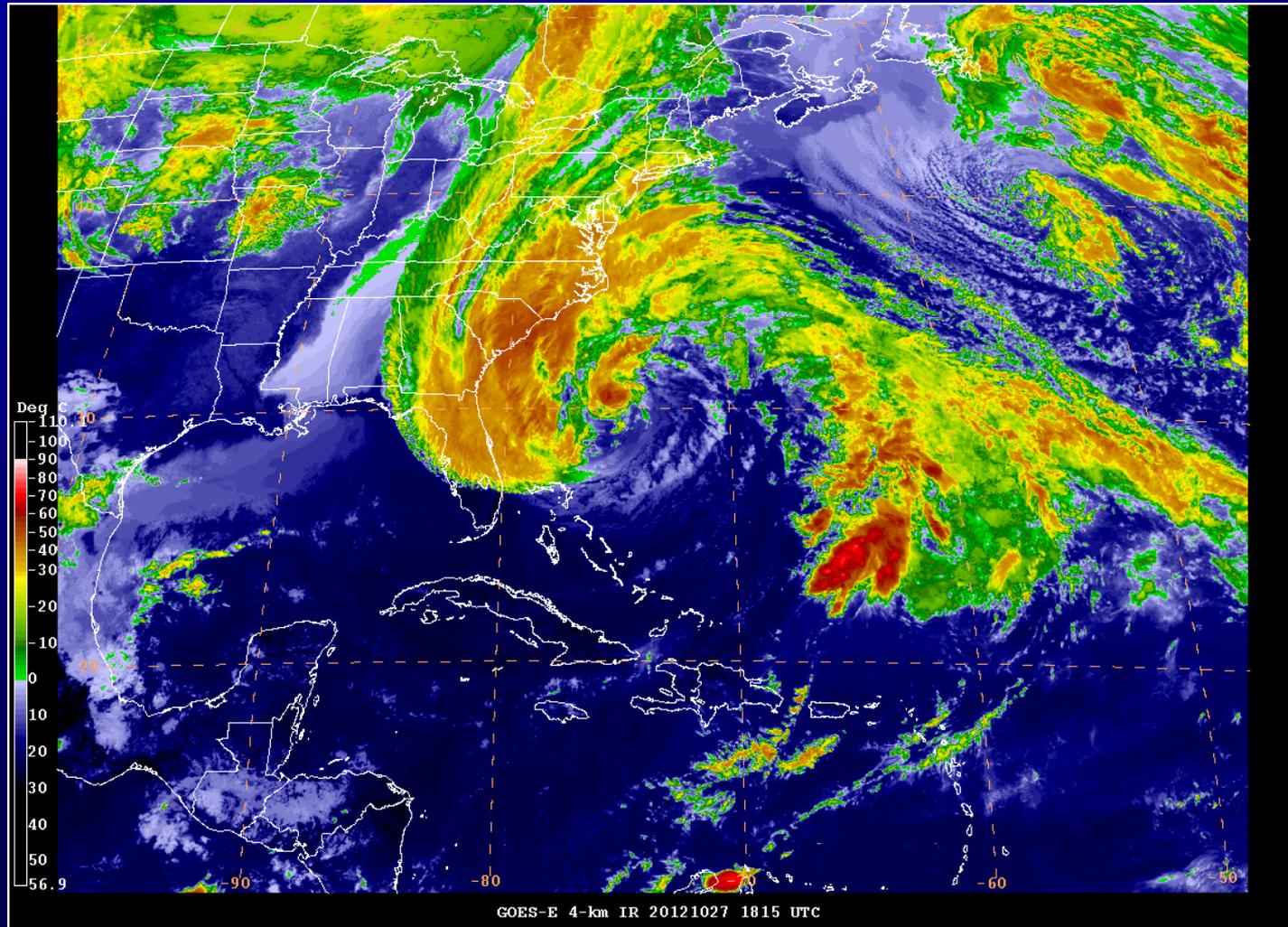
34 kt wind radii:

NE 450

SE 300

SW 240

NW 280





Hurricane Sandy (2012)

00 UTC 8 October

Intensity 65 kt

Pressure 960 mb

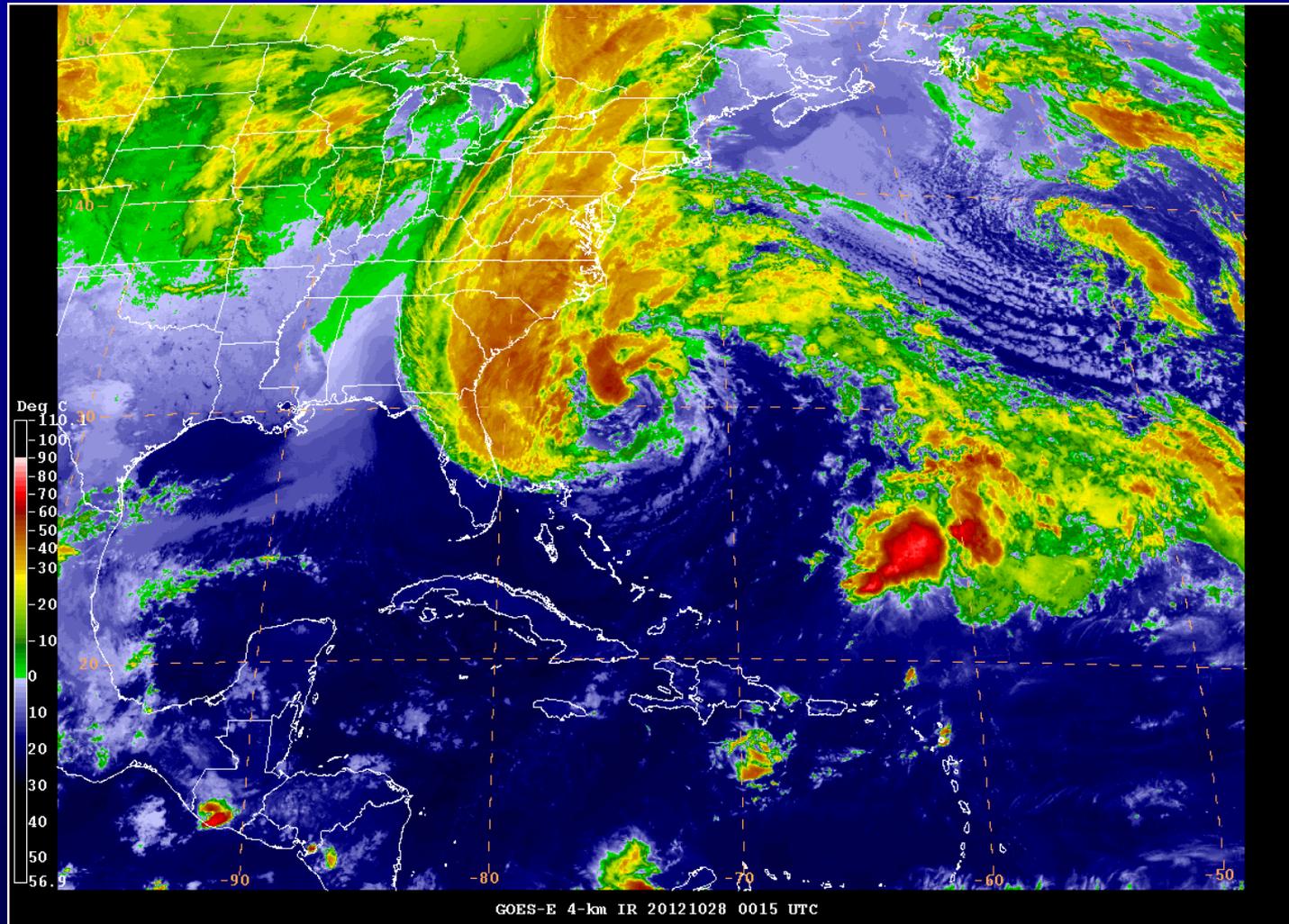
34 kt wind radii:

NE 480

SE 300

SW 300

NW 280





Hurricane Sandy (2012)



06 UTC 8 October

Intensity 65 kt

Pressure 959 mb

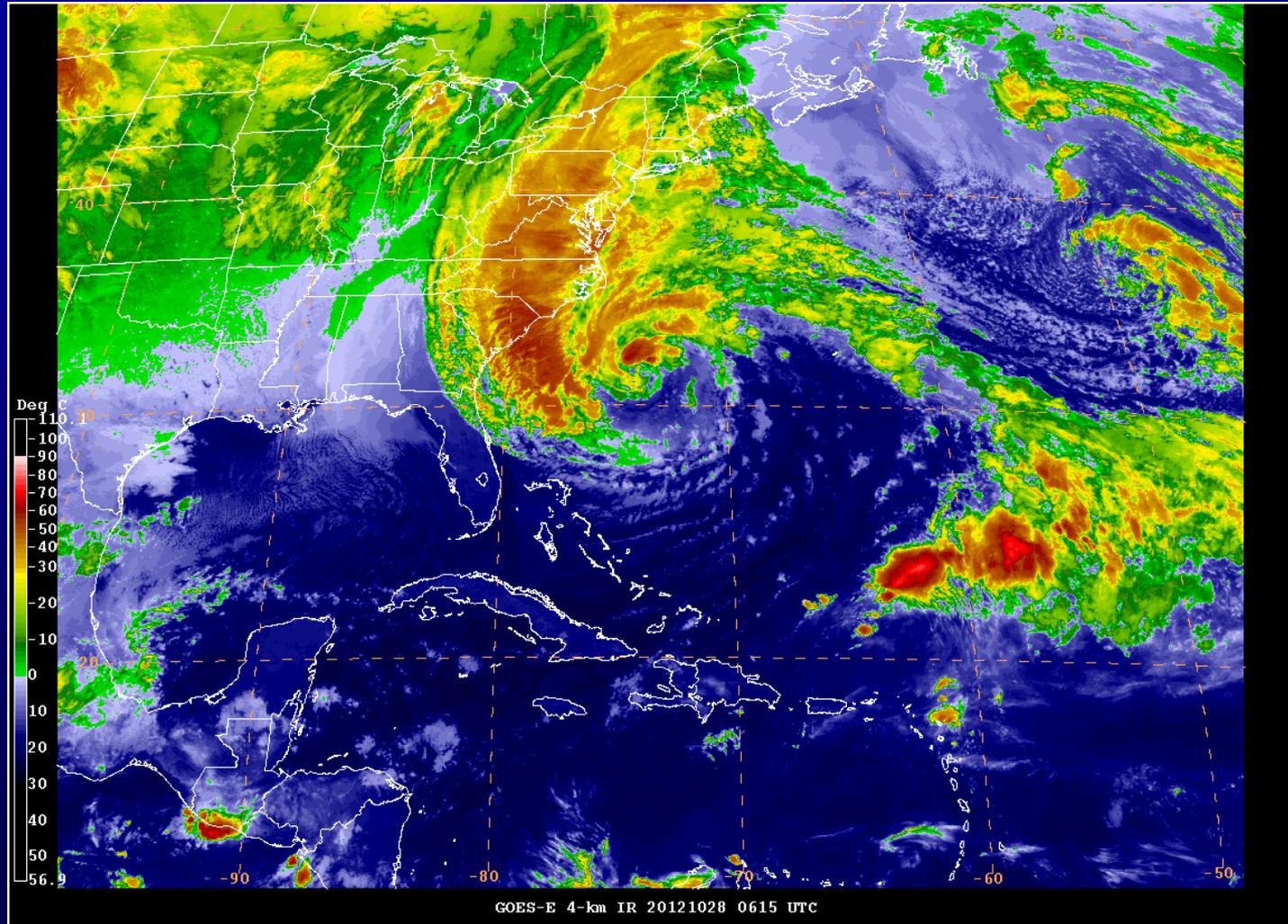
34 kt wind radii:

NE 450

SE 300

SW 300

NW 270





Hurricane Sandy (2012)

12 UTC 8 October

Intensity 65 kt

Pressure 954 mb

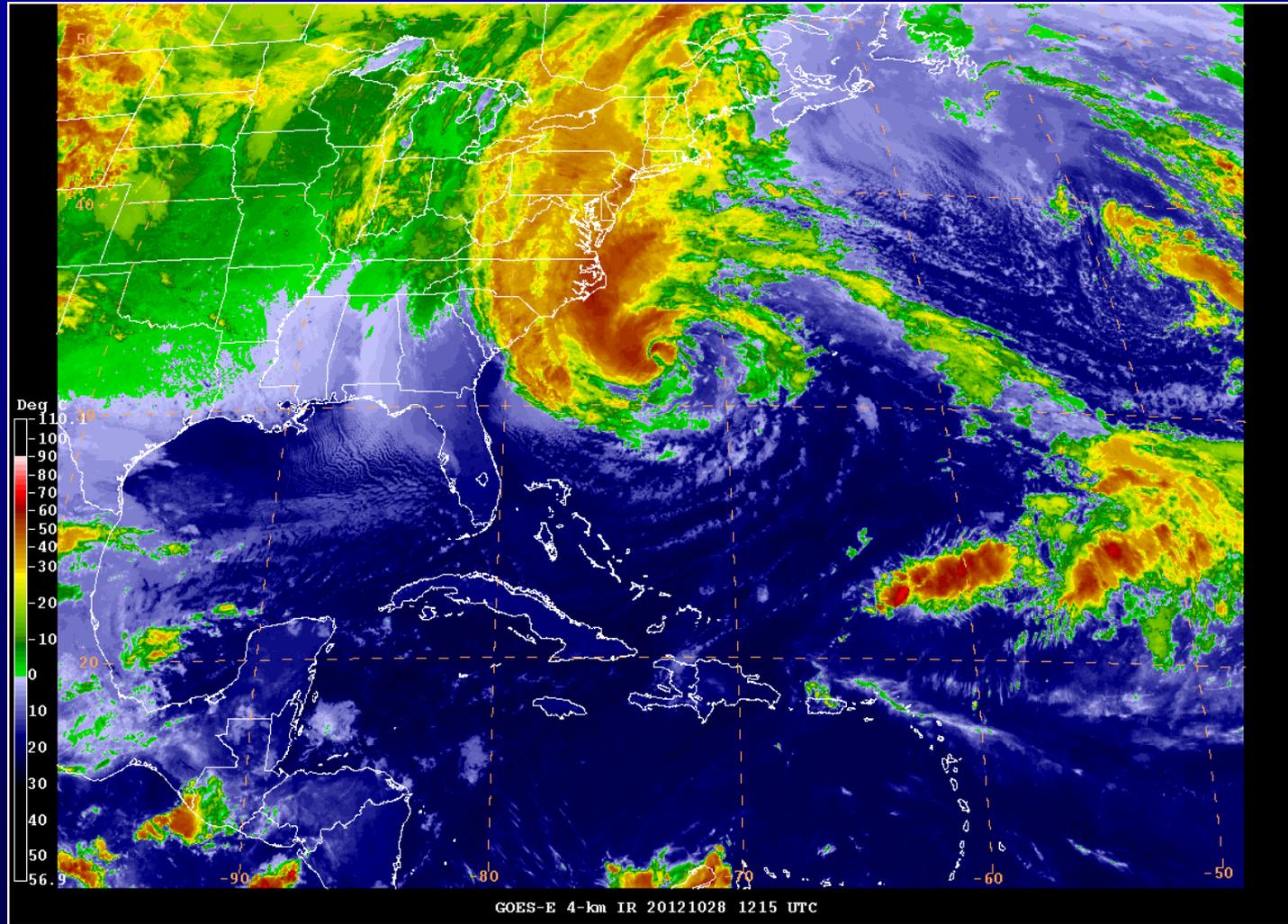
34 kt wind radii:

NE 450

SE 300

SW 300

NW 270





Hurricane Sandy (2012)



18 UTC 8 October

Intensity 65 kt

Pressure 952 mb

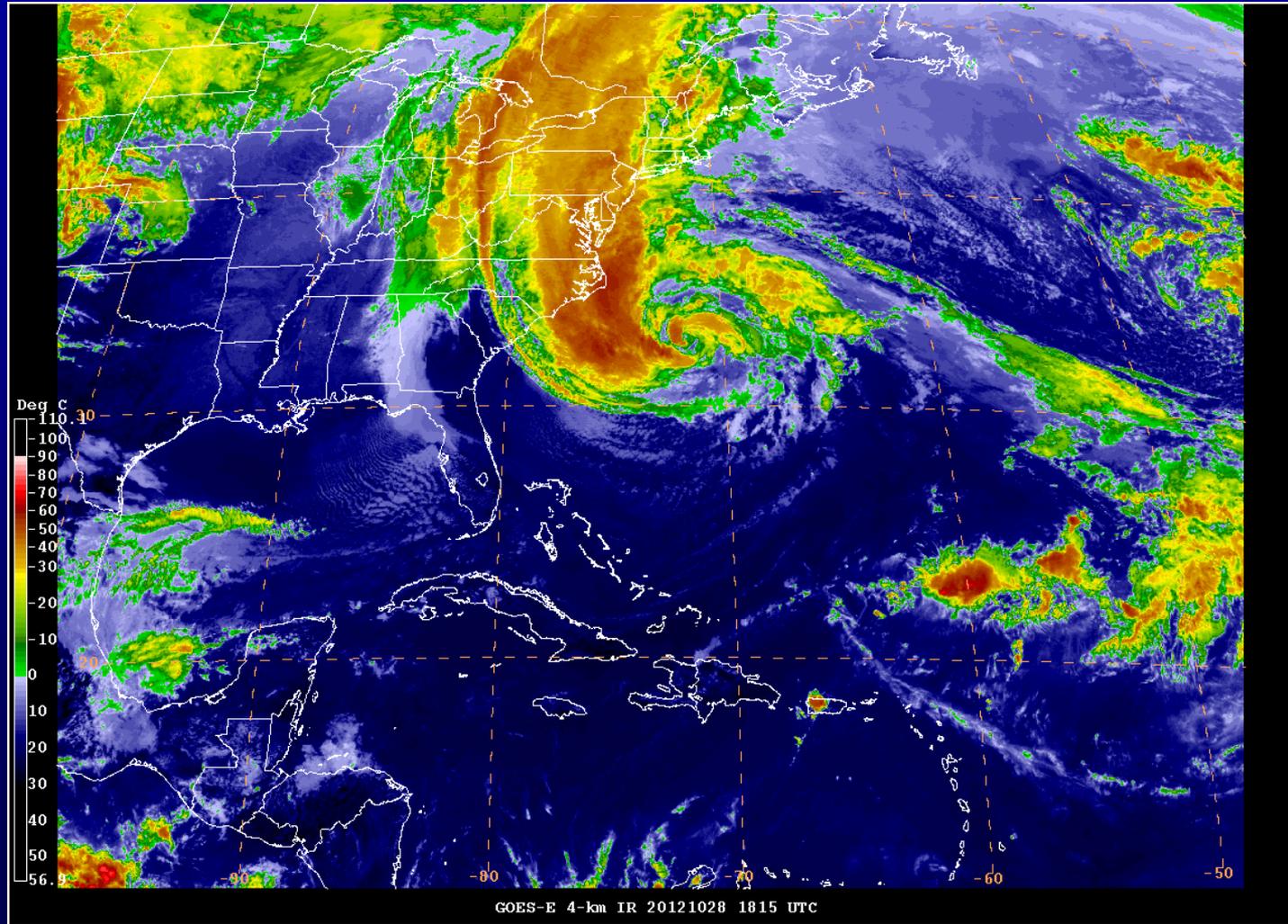
34 kt wind radii:

NE 450

SE 300

SW 350

NW 270





Hurricane Sandy (2012)

00 UTC 9 October

Intensity 70 kt

Pressure 950 mb

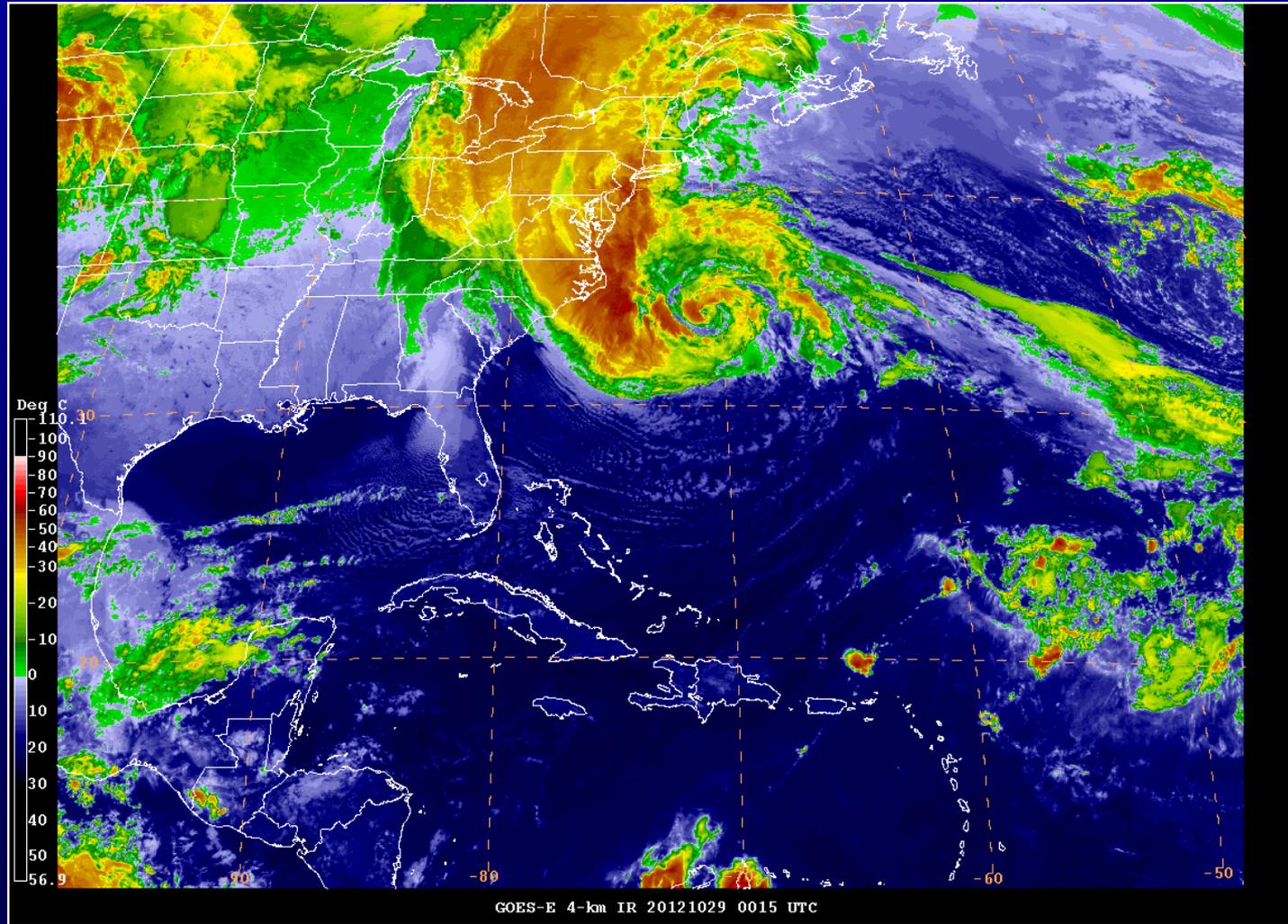
34 kt wind radii:

NE 450

SE 300

SW 400

NW 270





Hurricane Sandy (2012)

06 UTC 9 October

Intensity 80 kt

Pressure 947 mb

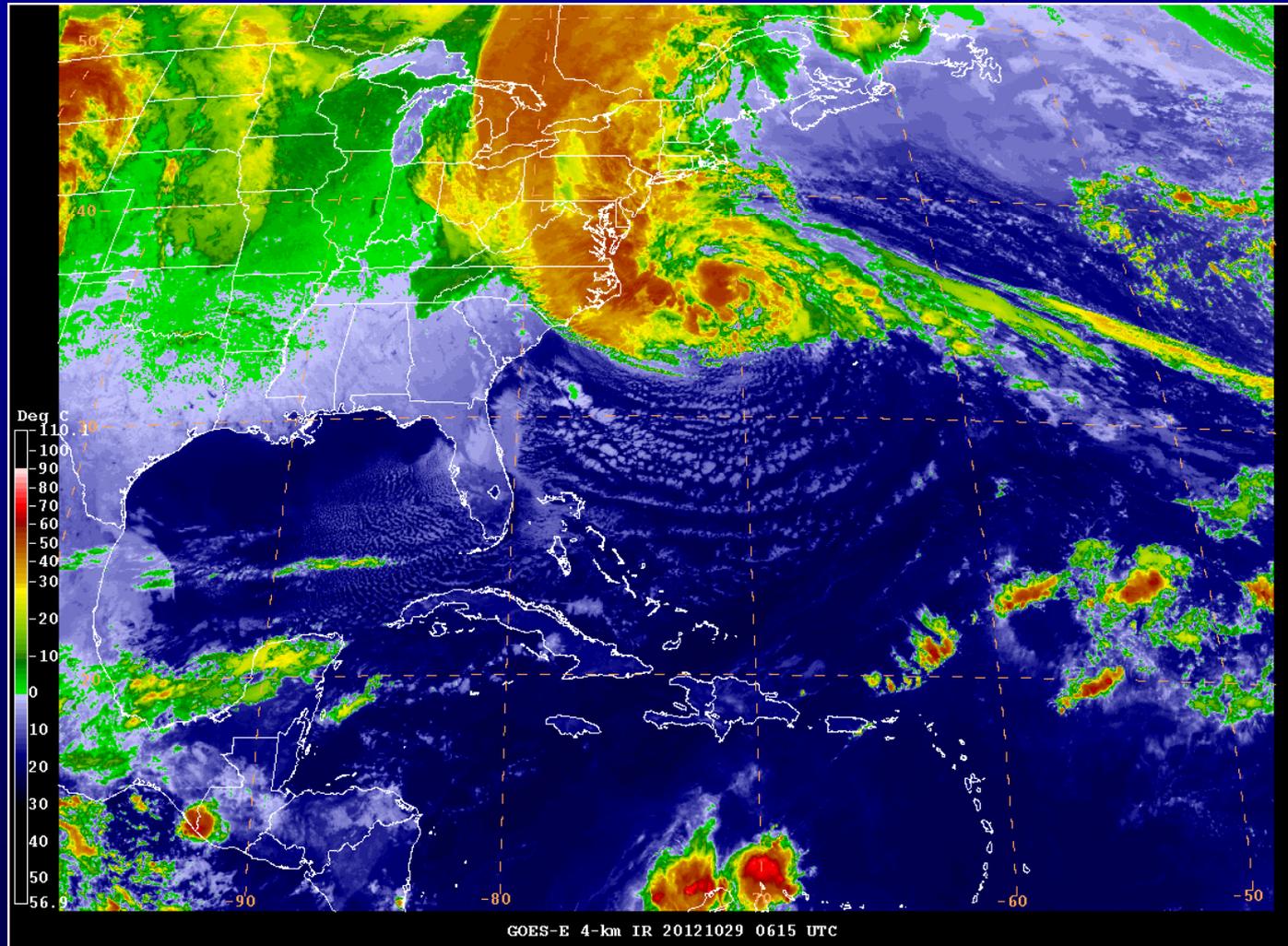
34 kt wind radii:

NE 420

SE 360

SW 450

NW 270



GOES-E 4-km IR 20121029 0615 UTC



Hurricane Sandy (2012)

12 UTC 9 October

Intensity 85 kt

Pressure 945 mb

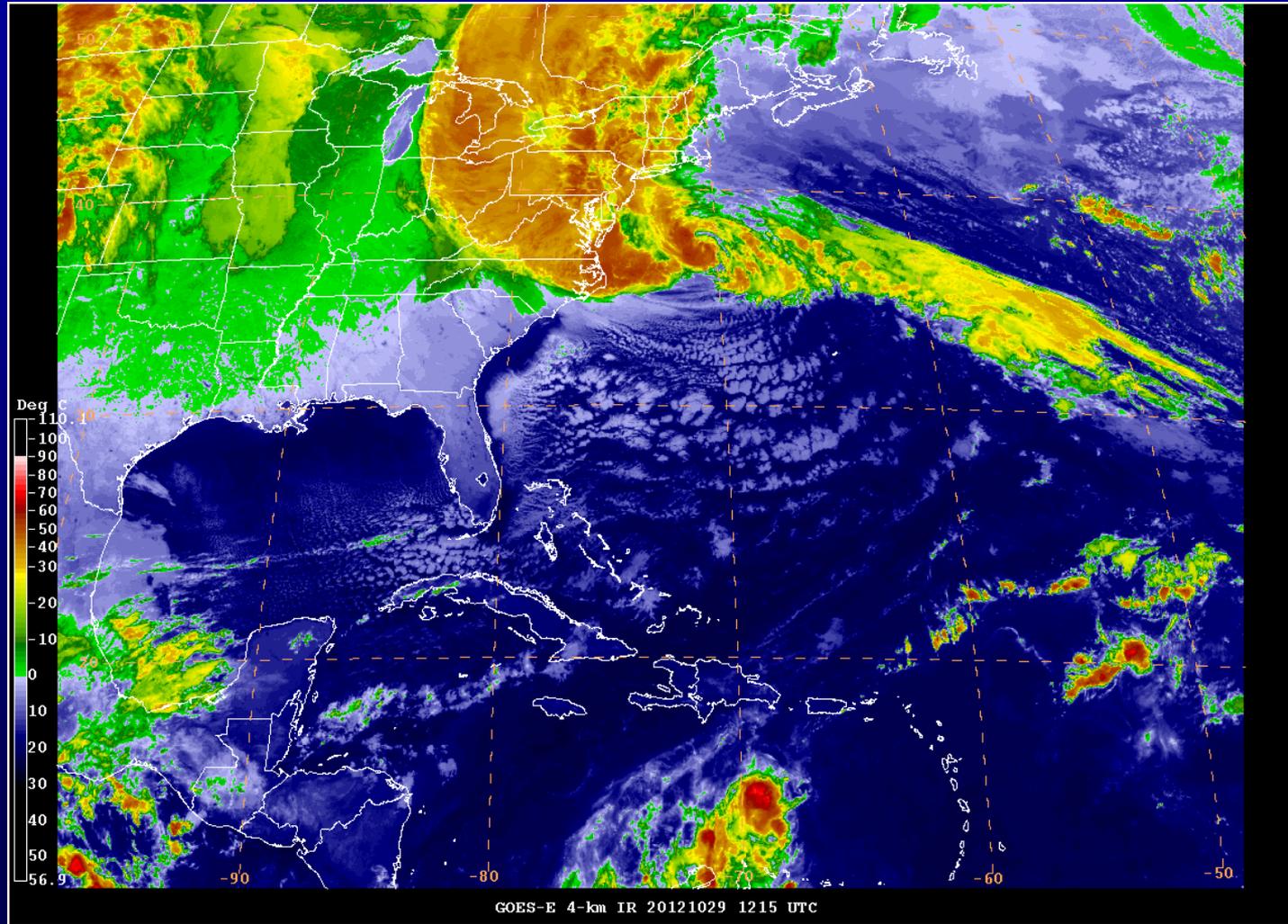
34 kt wind radii:

NE 420

SE 420

SW 400

NW 270





Hurricane Sandy (2012)

18 UTC 9 October

Intensity 80 kt

Pressure 940 mb

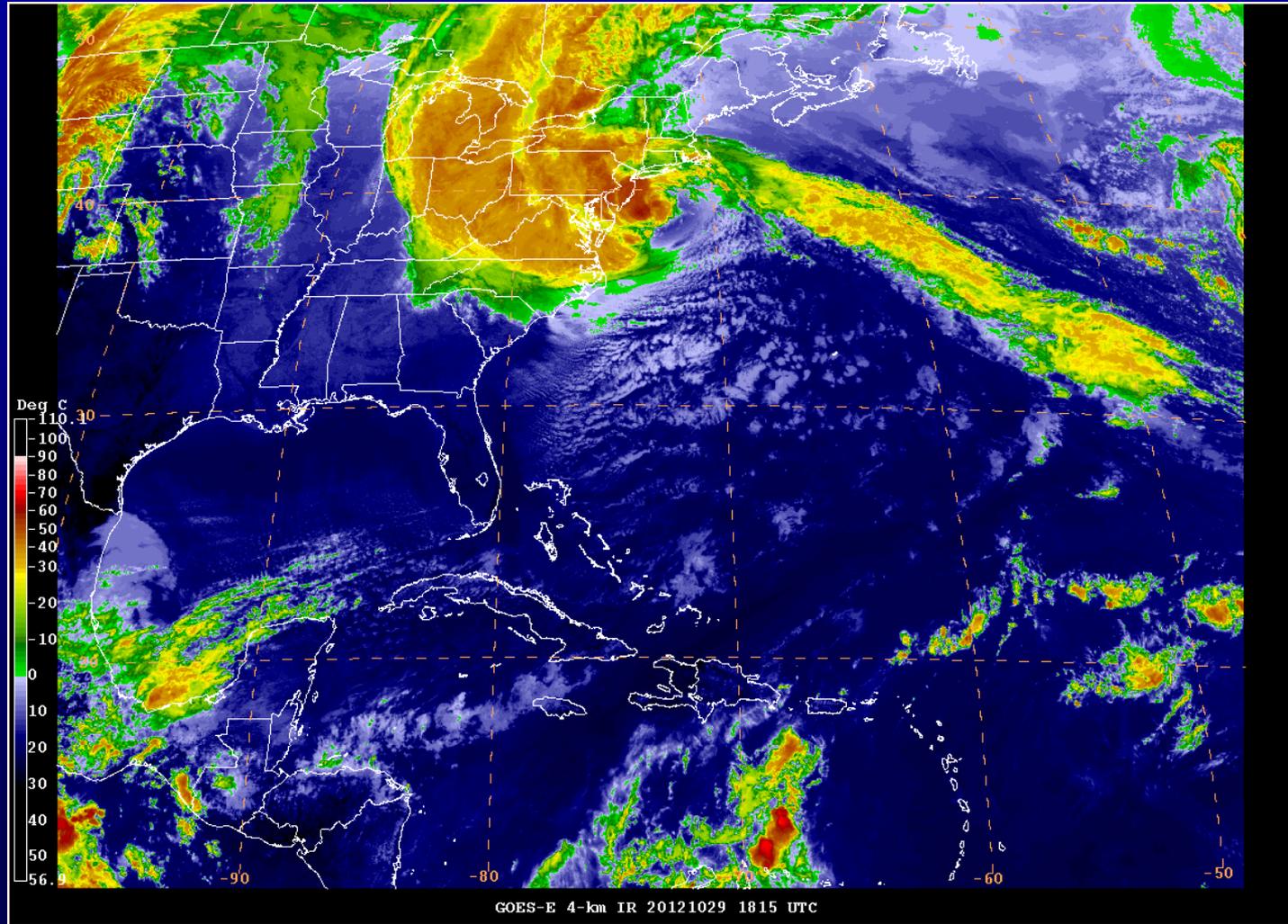
34 kt wind radii:

NE 420

SE 420

SW 400

NW 420





Hurricane Sandy (2012)

00 UTC 10 October

Intensity 70 kt

Pressure 946 mb

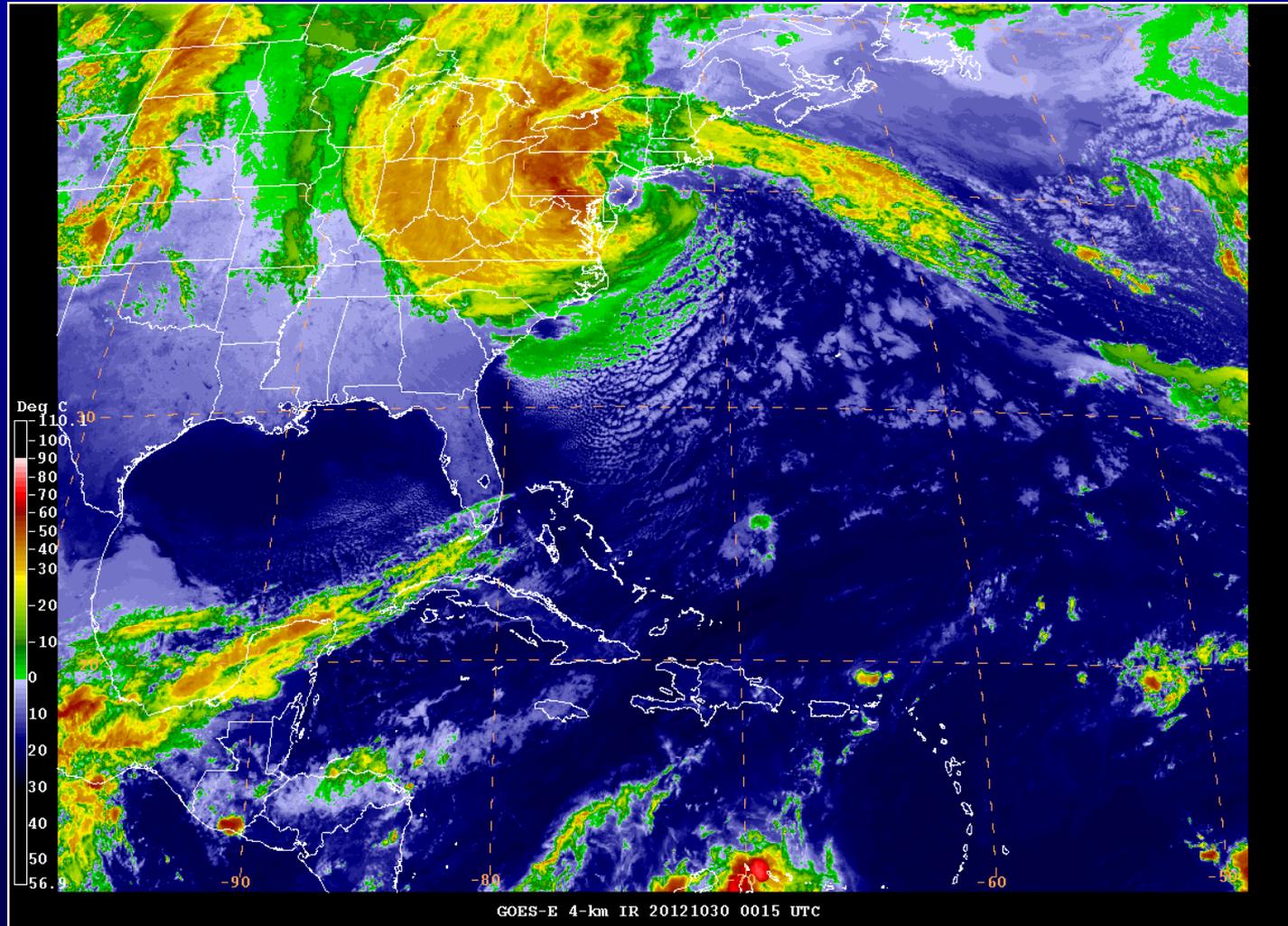
34 kt wind radii:

NE 460

SE 370

SW 400

NW 490



Extratropical



Hurricane Sandy (2012)

06 UTC 10 October

Intensity 55 kt

Pressure 960 mb

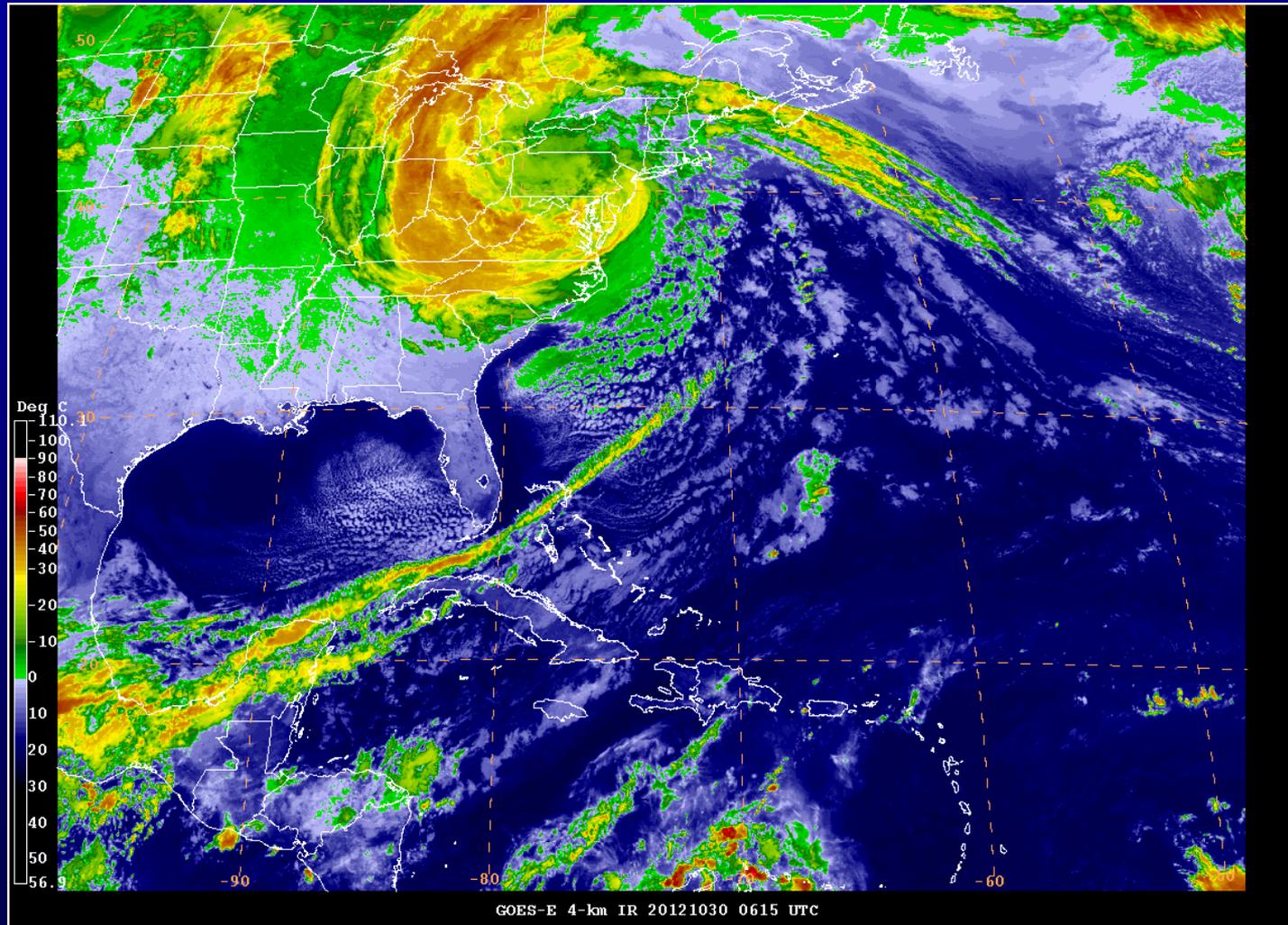
34 kt wind radii:

NE 450

SE 400

SW 160

NW 530



Extratropical



Hurricane Sandy (2012)

12 UTC 10 October

Intensity 50 kt

Pressure 978 mb

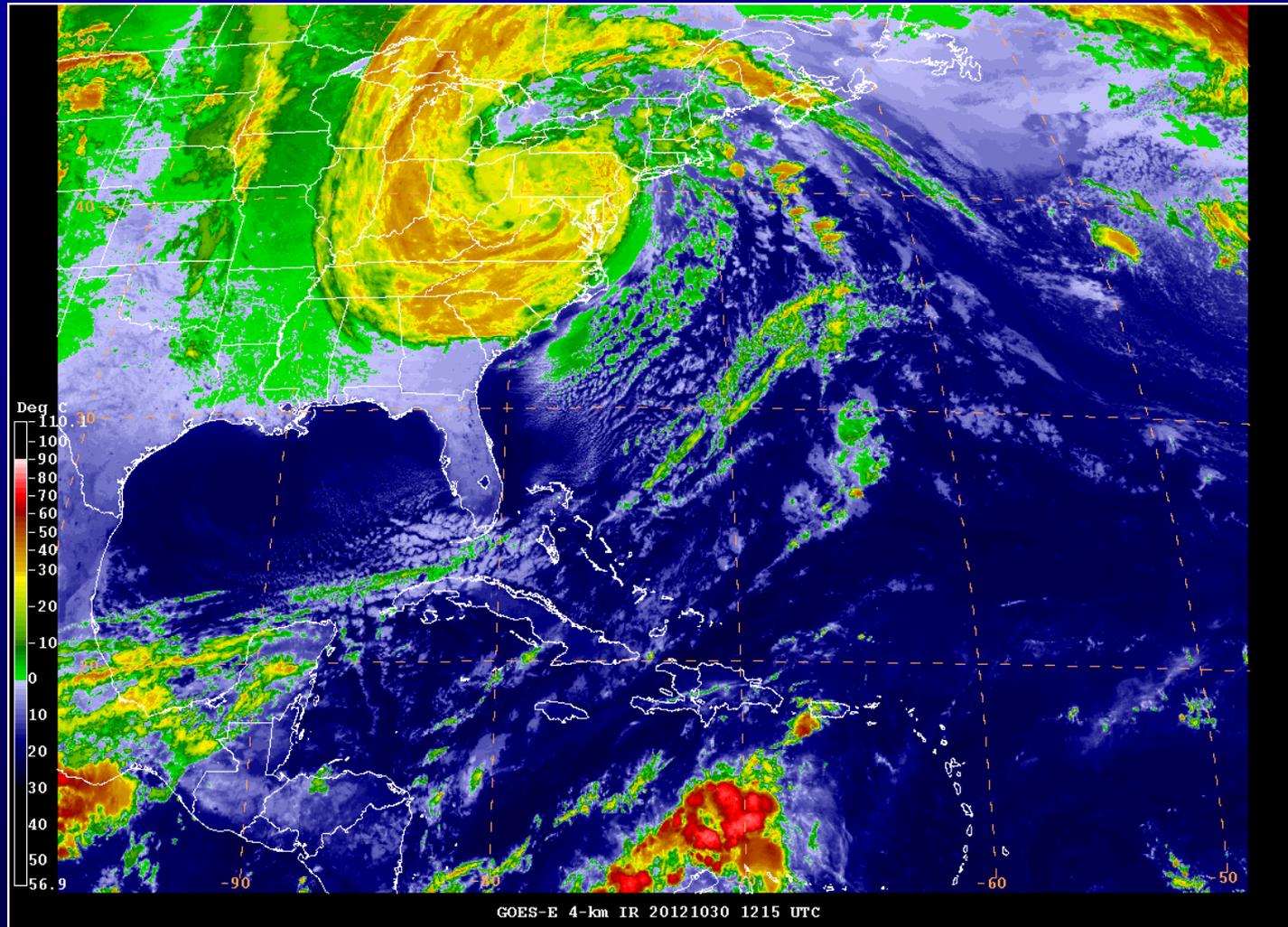
34 kt wind radii:

NE 450

SE 490

SW 0

NW 500



Extratropical



Environmental and Structural Changes that Occur during ET



Environmental Changes Encountered During ET

(Jones et al. 2003)

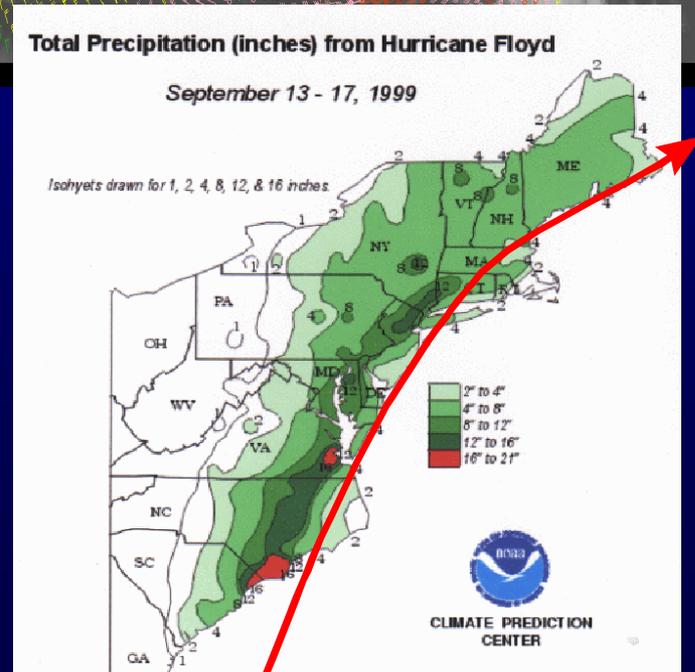
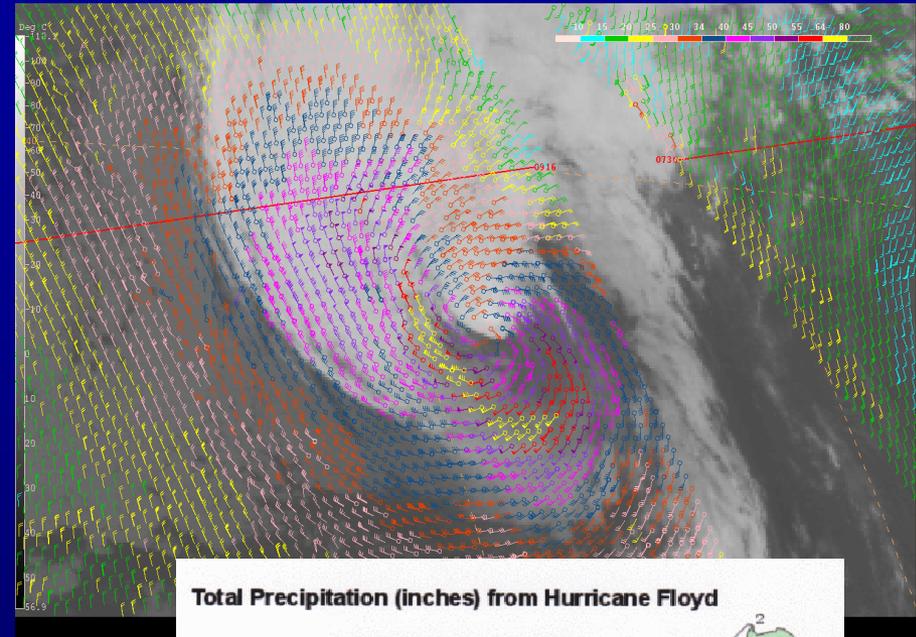


- Decreasing SST as TC moves poleward
 - Drier/more stable lower-tropospheric airmass
 - Distribution of sensible and latent heat fluxes from ocean surface becomes more asymmetric and variable
- Interaction with large-scale baroclinic features
 - Upper-level troughs/jets
 - Lower-tropospheric fronts
 - SST gradients
 - Extratropical cyclones
- Increased wind shear
- Landfall or interaction with continental landmasses

Structural Changes Observed During ET

(Jones et al. 2003)

- Primary energy source transitions from latent heat release to baroclinic processes
 - Decrease in deep convection near center
 - Thermal advection, synoptic forcing (CVA, jet dynamics) become sources for energy
 - Both can be important during ET
- Impact of surface fluxes from the ocean surface change
 - Crucial to TC development, but more varied impact on extratropical cyclones
- Increase in baroclinity across cyclone
 - Development of fronts, asymmetric cloud and precipitation patterns due to strong thermal advectons
- Reduction in maximum wind (usually but not always)
- Expansion of outer wind field





Conceptual Models and Composites of ET



Satellite-Imagery Based Conceptual Model

Klein et al. (2000)



- Examined satellite imagery and NWP analyses for 30 cyclones undergoing ET in the Northwest Pacific 1994–1998
- Three step transformation stage
 - Begins when cyclone starts to interact with baroclinic zone
 - Satellite imagery shows developing asymmetry in clouds and large decrease of deep convection on western side of TC
 - Ends when cyclone is fully embedded in baroclinic zone
 - Satellite imagery shows baroclinic features and cyclone center embedded in cold, descending air

Common Satellite Features During ET

Klein et al. (2000)

Decrease in cloudiness in western semicircle

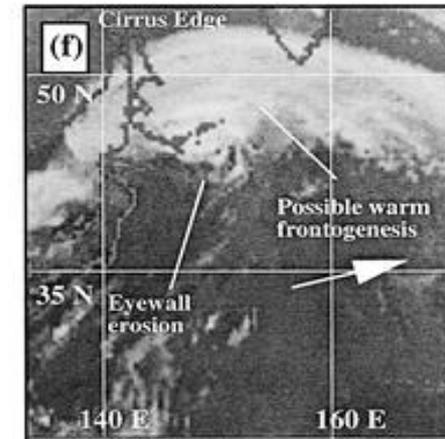
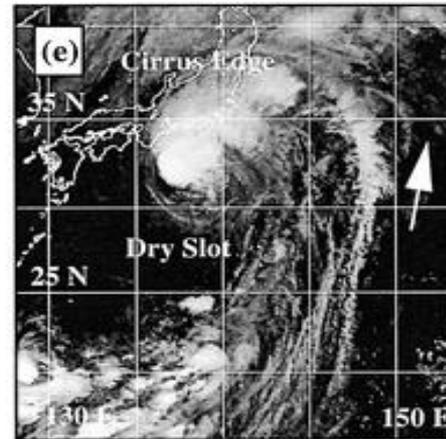
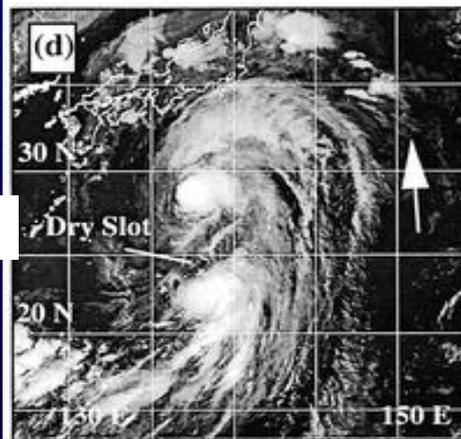
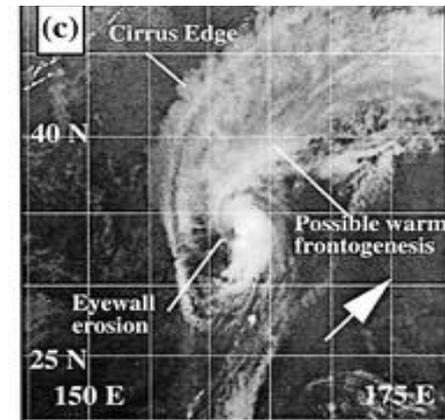
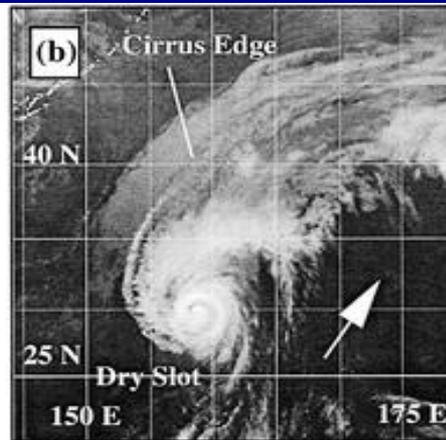
Development of dry slot

Convective asymmetry

Cirrus shield marks interaction with polar jet

Inner core deep convection eroded – gap in eyewall

Signals of strong warm frontogenesis, only weak cold frontogenesis



Step 1

Step 2

Step 3

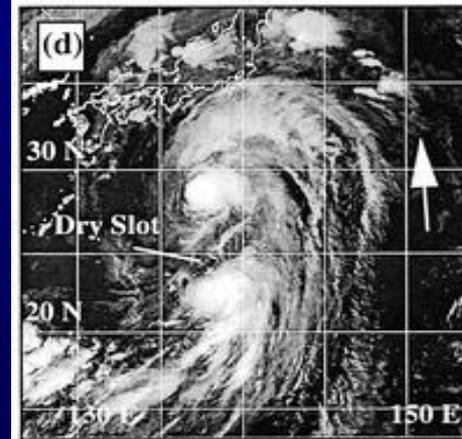
Supertyphoon
Ginger

Typhoon Opal

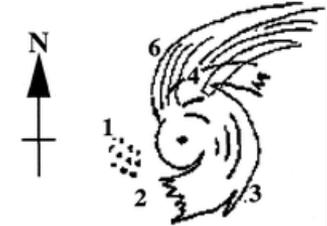
NW Pacific ET Conceptual Model

Klein et al. (2000)

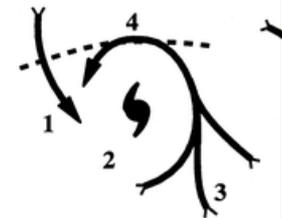
- Step 1 – Transformation Begins
 - Cyclone moves over cooler SSTs
 - Interaction begins with baroclinic zone
 - Cold, dry advection begins west of center
 - Convection decreases in western semicircle of outer circulation
 - Dry slot forms in southwest quadrant
 - Warm, moist advection maintains convection in northeast quadrant
 - Cirrus shield develops in poleward outflow



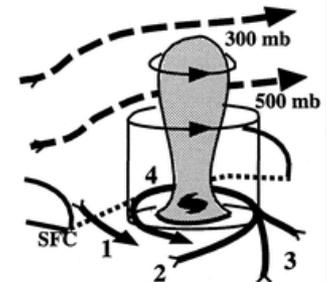
(a) Infrared Imagery



(b) Plan View (Storm Rela



(c) 3-D View (Storm Rela



--- = Po

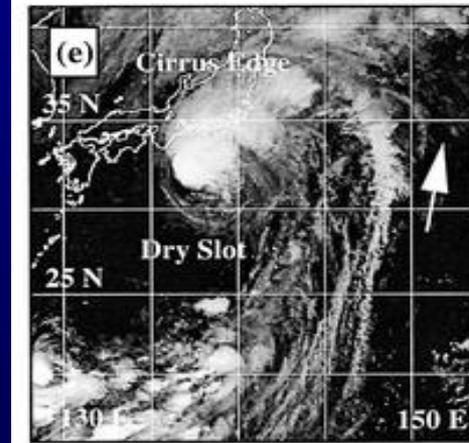
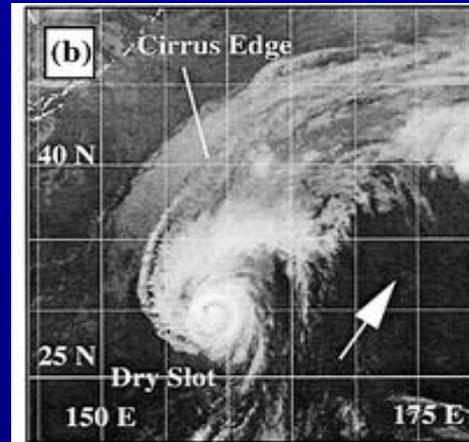
STEP 1

Klein et al. (2000)
Weather and Forecasting

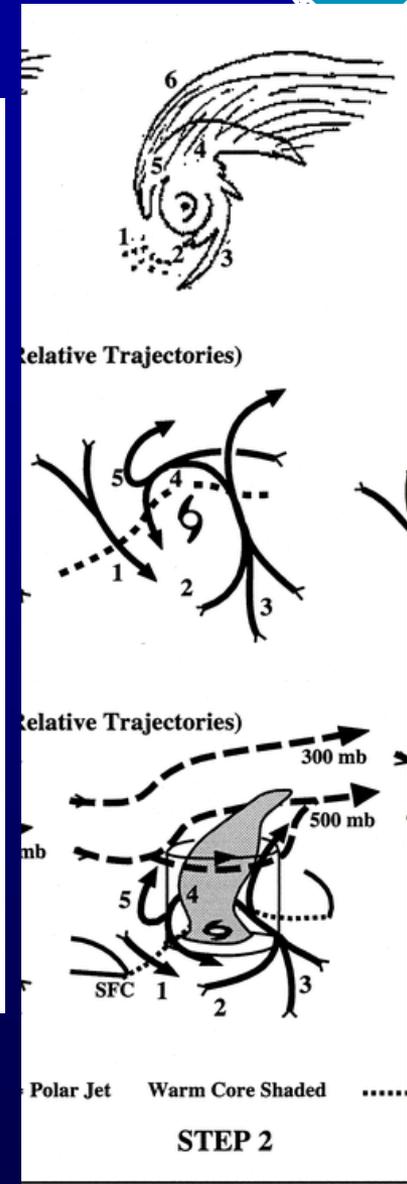
NW Pacific ET Conceptual Model

Klein et al. (2000)

- Step 2 – Transformation Continues
 - Cyclone just south of baroclinic zone
 - Thermal advection increases as cyclone circulation impinges on baroclinic zone
 - Cloud pattern asymmetry increases
 - Dry slot grows in size
 - Increasing vertical wind shear begins to distort vertical alignment of TC inner core at upper levels
 - Convection persists in inner core



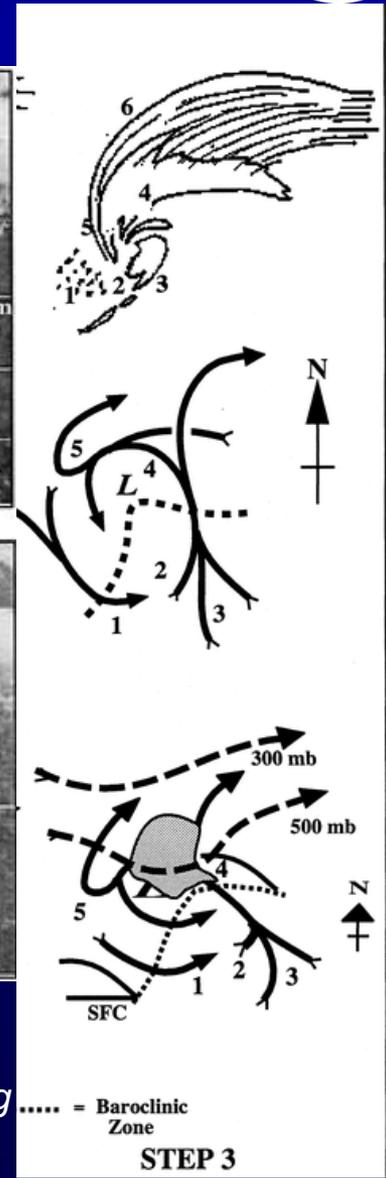
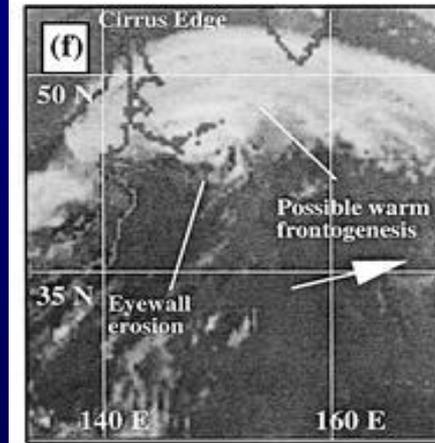
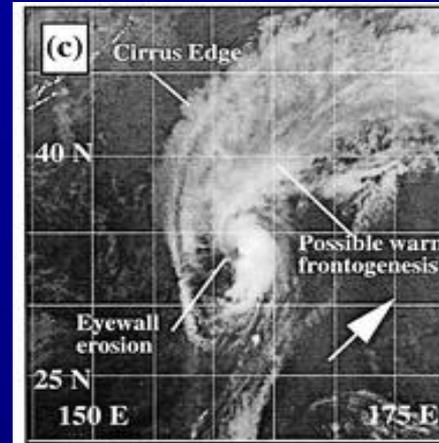
Klein et al. (2000)
Weather and Forecasting



NW Pacific ET Conceptual Model

Klein et al. (2000)

- Step 3 – Transformation Complete
 - Cyclone completely embedded in baroclinic zone
 - Thermal advections continue to intensify
 - Cloud asymmetry grows
 - Vertical shear advects upper-level warm core downstream
 - Weaker, lower-level warm core persists over cyclone center
 - Inner-core convection disrupted and eyewall erodes on south and west sides
 - Pronounced warm frontal cloud band and weaker cold frontal cloud band visible



Klein et al. (2000)
Weather and Forecasting



North Atlantic ET Study

Hart et al. (2006)



- Based on 34 Atlantic ET events from 1998–2003 examined using 1° NOGAPS analysis
- Examined characteristics that differed between cyclones for
 - Fast vs. slow ET
 - Post ET intensification vs. weakening
 - Post ET structure – cold core or warm seclusion



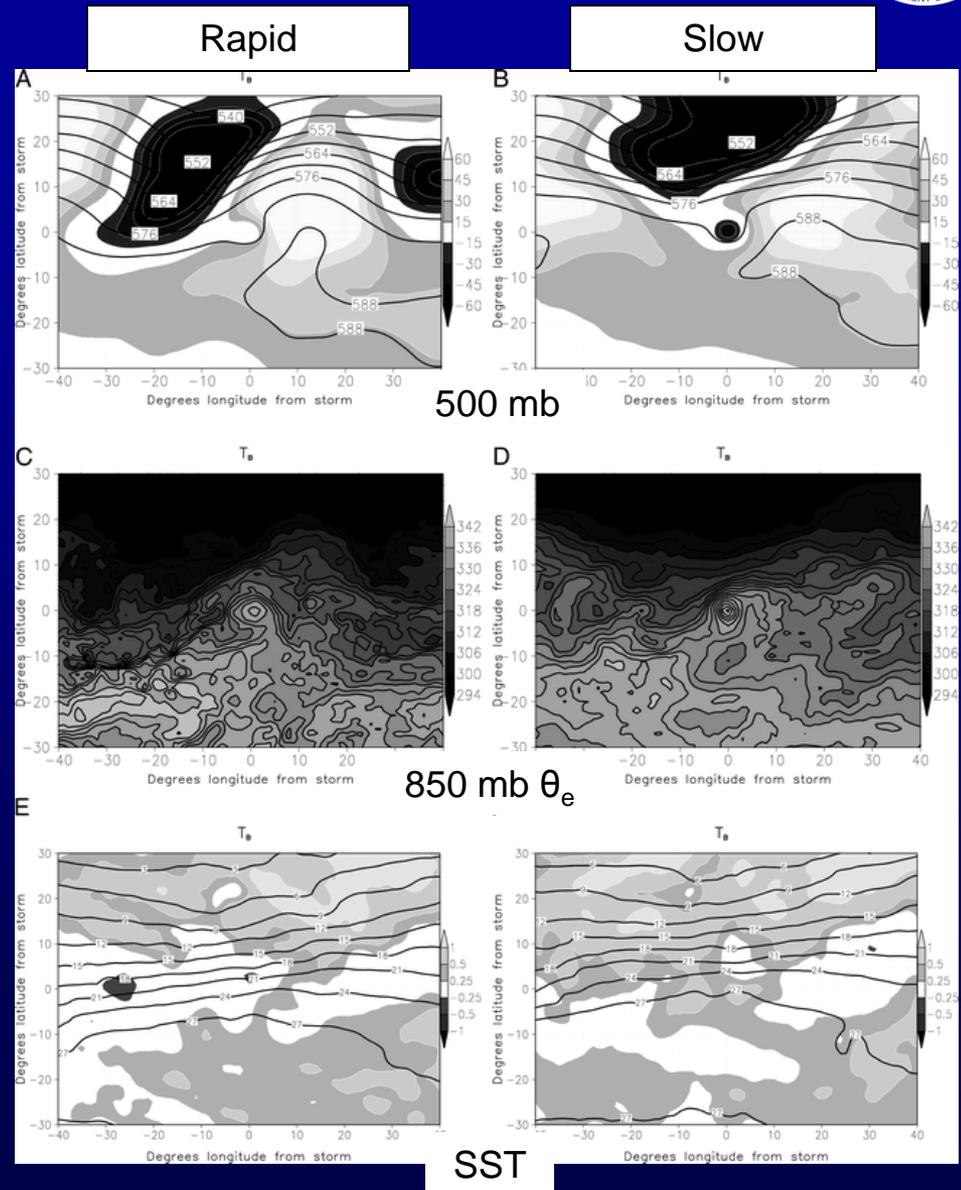
North Atlantic ET – Slow or Fast Transition?

Hart et al. (2006)



Rapid or slow transition?

- **Rapid ET (≤ 12 h)**
 - Longwave pattern meridional
 - TC initially smaller and weaker than average
 - SST $\ll 27^\circ\text{C}$, with large SST gradient
- **Slow ET (≥ 48 h, extended hybrid phase)**
 - Longwave pattern zonal
 - SST $\approx 27^\circ\text{C}$
 - TC initially stronger and larger
 - Environmental forcing still supports tropical development



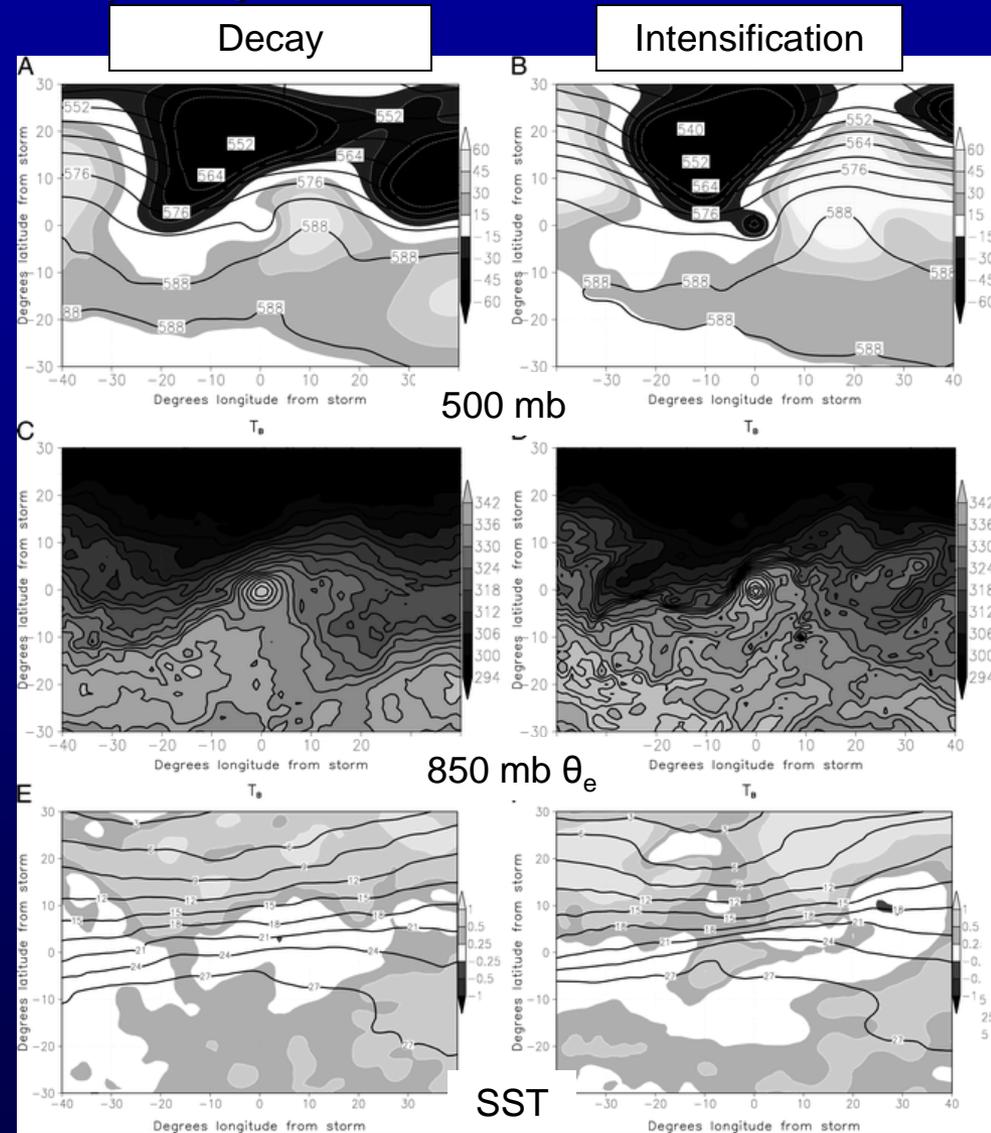


North Atlantic ET – Post Transition Intensification or Decay? Hart et al. (2006)



Post-transition intensification or decay?

- **Decay**
 - Positively-tilted upper trough
 - Remnant TC weak and detached from tropical air
 - Warmer SST but weaker gradient
- **Intensification**
 - Negatively-tilted upper trough
 - Remnant TC intense and more connected to tropics
 - Stronger SST gradient





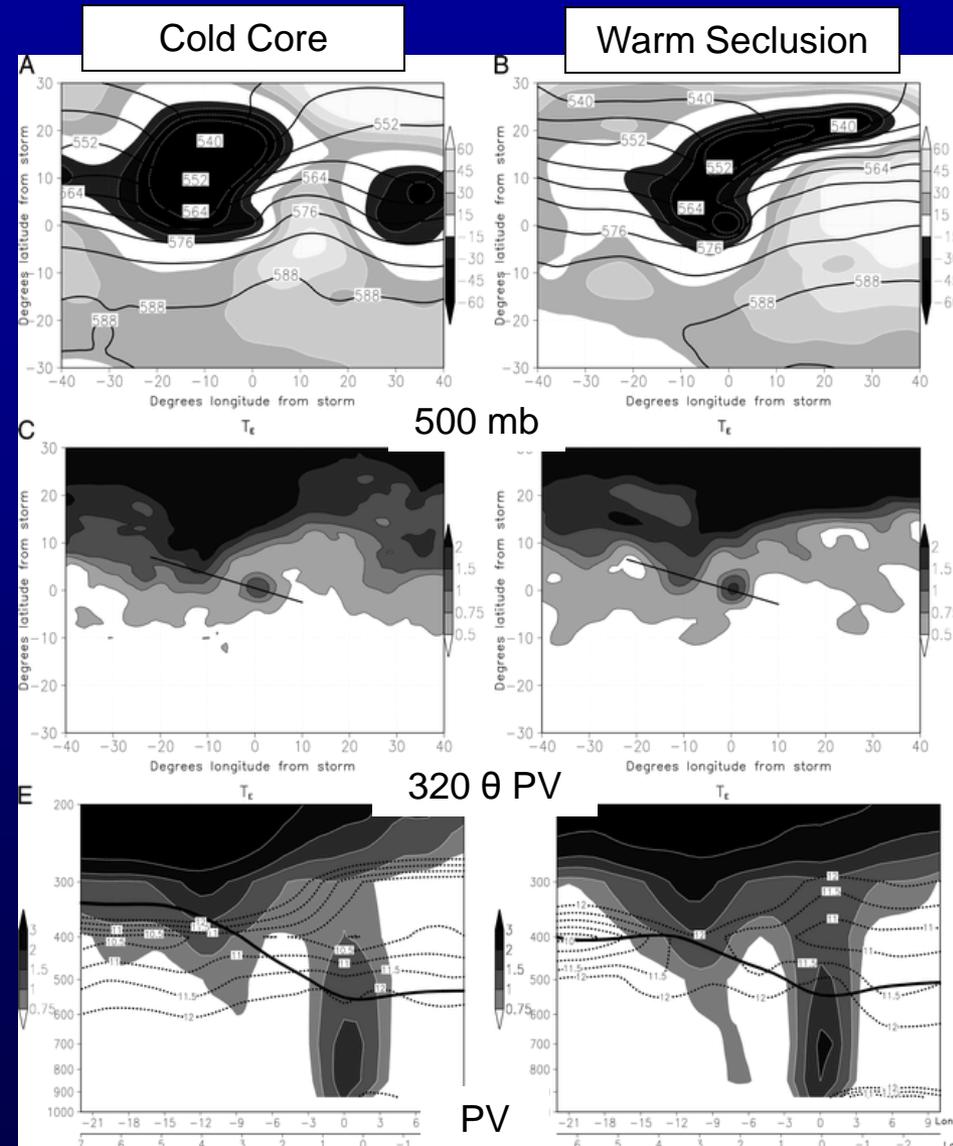
North Atlantic ET – Post Transition Structure

Hart et al. (2006)



Post-transition structure

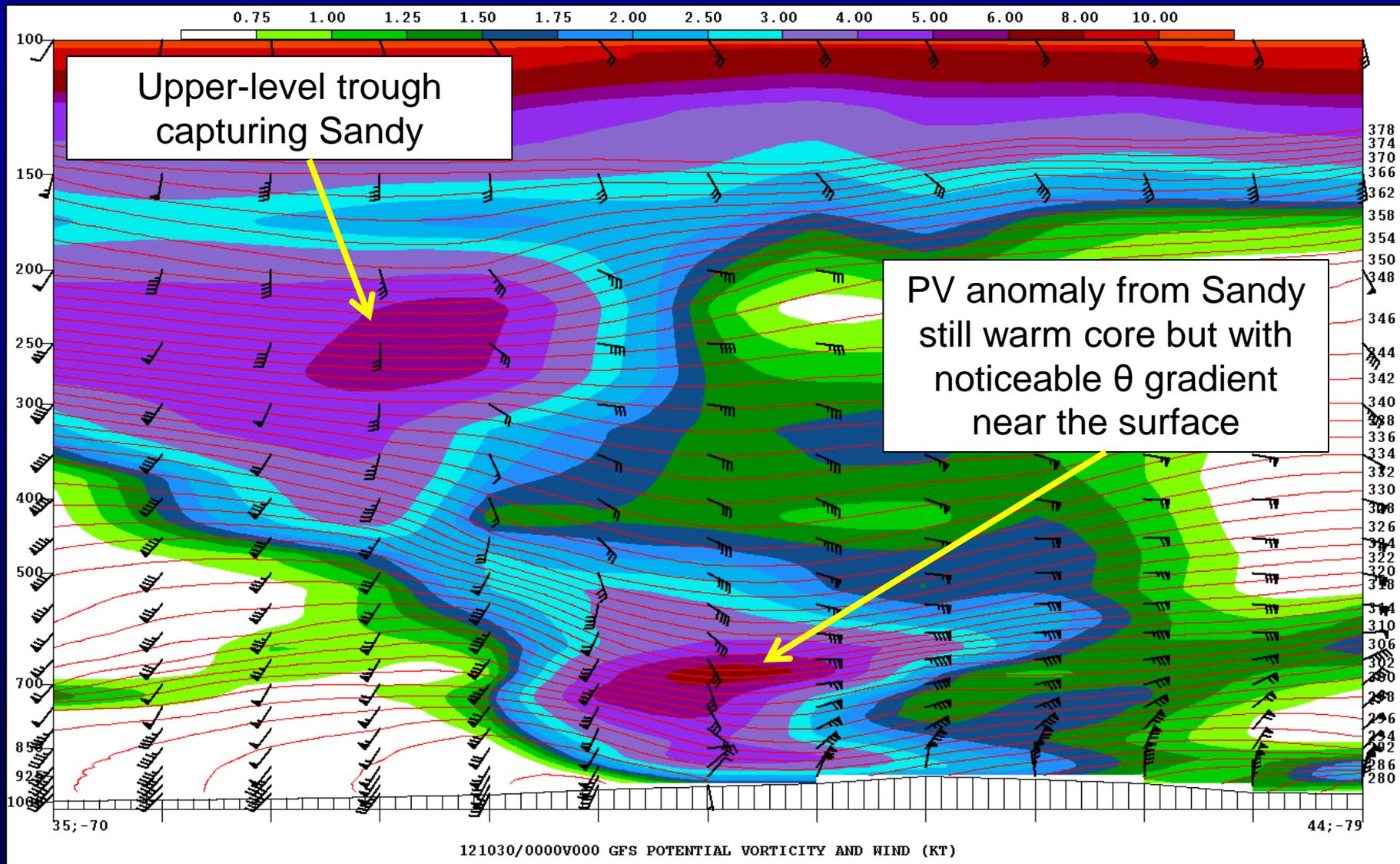
- **Cold Core**
 - Large horizontal scale of upper trough does not match scale of TC
 - Trough has smaller vertical extent
- **Warm Seclusion**
 - Scale of upper trough closely matches that of TC (usually larger TCs)
 - Narrow horizontal extent of trough, but deep vertical extent
 - Increased wind threat due to expansion of gale force wind radii *and* strong maximum wind
 - More difficult for models to forecast





Sandy Post-Transition

GFS Cross Section of PV, θ , Wind 00Z 30 October 2012





Forecasting Tools





Phase Space Diagrams

Hart (2003)



- Cyclone “phase space” diagrams provide a way to visualize the location of a cyclone in the continuum between tropical and extratropical
- Constructed from model analyses and forecasts for active cyclones – available online at: <http://moe.met.fsu.edu/cyclonephase/>
- Useful for determining character of model analyzed cyclone and forecast trends
- Three parameters used to represent the “phase” of a cyclone:
 1. B – Storm-motion-relative 900–600-mb thickness gradient across cyclone
 2. $-V_T^L$ – Magnitude of lower-tropospheric (900–600 mb) cyclone thermal wind
 3. $-V_T^U$ – Magnitude of upper-tropospheric (600–300 mb) cyclone thermal wind

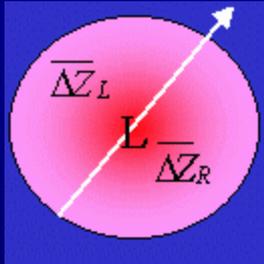
Two phase diagrams constructed plotting 1 vs. 2 and 2 vs. 3



Storm-Motion-Relative Thickness Gradient (B)

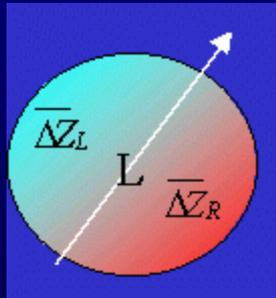


- Measures the degree of thermal asymmetry across a cyclone relative to the direction of motion
- Provides a representation of how “frontal” a cyclone is



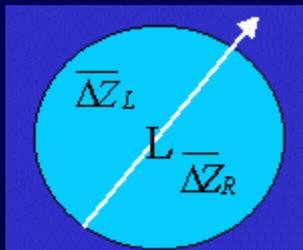
Mature tropical cyclone $\rightarrow B = 0 \rightarrow$ Symmetric

Warmest temperature at center decreasing uniformly in all directions



Developing or mature extratropical cyclone $\rightarrow B \gg 0 \rightarrow$ Asymmetric

Pronounced thermal gradient across cyclone with well defined “warm” and “cold” sectors



Occluded extratropical cyclone $\rightarrow B = 0 \rightarrow$ Symmetric

Uniform thermal structure with little change in temperature near cyclone center



Magnitude of Thermal Wind

- Lower Troposphere ($-V_T^L$)
 - Shows structure of the cyclonic circulation in the lower troposphere (900 to 600 mb)
- Middle-Upper Troposphere ($-V_T^U$)
 - Shows structure of the cyclonic circulation in the mid-to upper-troposphere (600 to 300 mb)

Circulation character

- **Warm core** ($-V_T > 0$) → circulation strength *decreases* with height
- **Cold core** ($-V_T < 0$) → circulation strength *increases* with height
- **Neutral** ($-V_T \approx 0$) → circulation strength shows little change with height



Phase Diagram 1



Thermal Asymmetry versus Lower-Tropospheric Thermal Wind

Symmetric warm core

- $B \leq 0$ and $-V_T^L > 0$
 - Tropical cyclones, warm seclusions

Asymmetric warm core

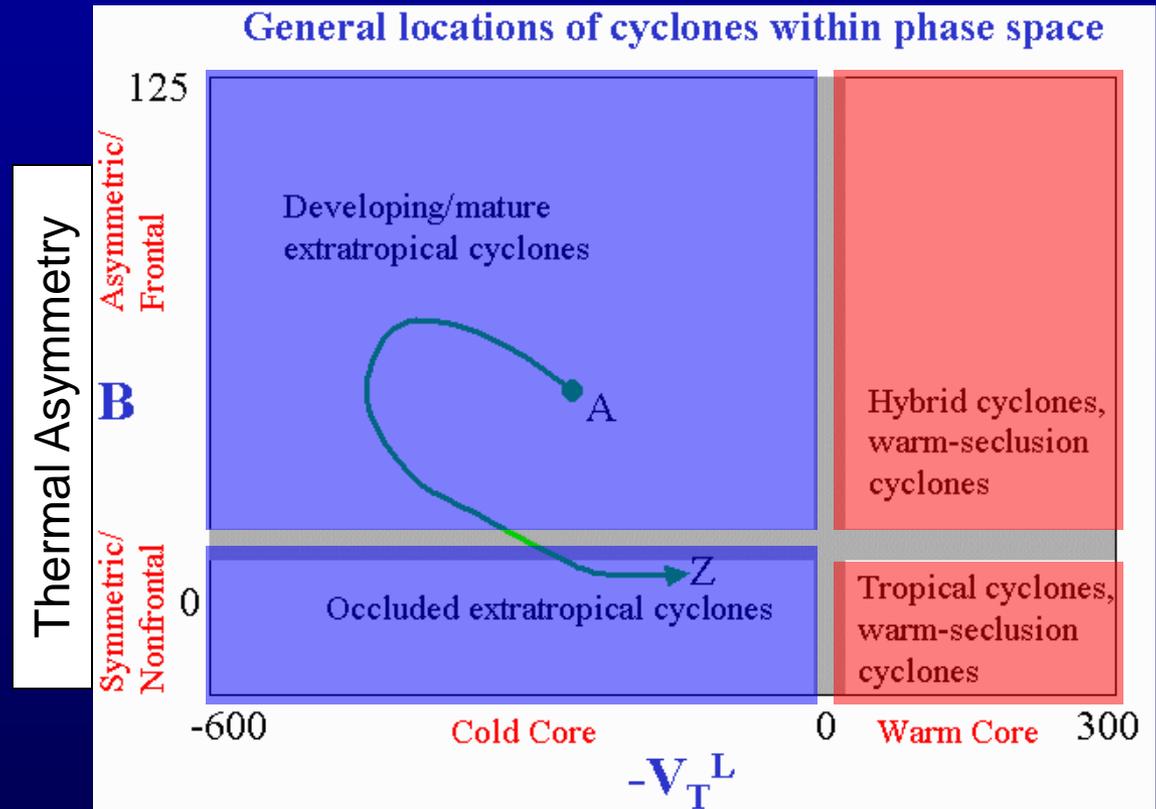
- $B > 0$ and $-V_T^L > 0$
 - Hybrid cyclones, warm seclusions
 - Most cyclones undergoing ET found here

Symmetric cold core

- $B \leq 0$ and $-V_T^L < 0$
 - Occluded extratropical cyclones

Asymmetric cold core

- $B > 0$ and $-V_T^L < 0$
 - Developing or mature extratropical cyclones



Lower Troposphere



Phase Diagram 2

Upper vs. Lower tropospheric Thermal Wind

Deep warm core

- $-V_T^L > 0, -V_T^U > 0$
- Tropical cyclones

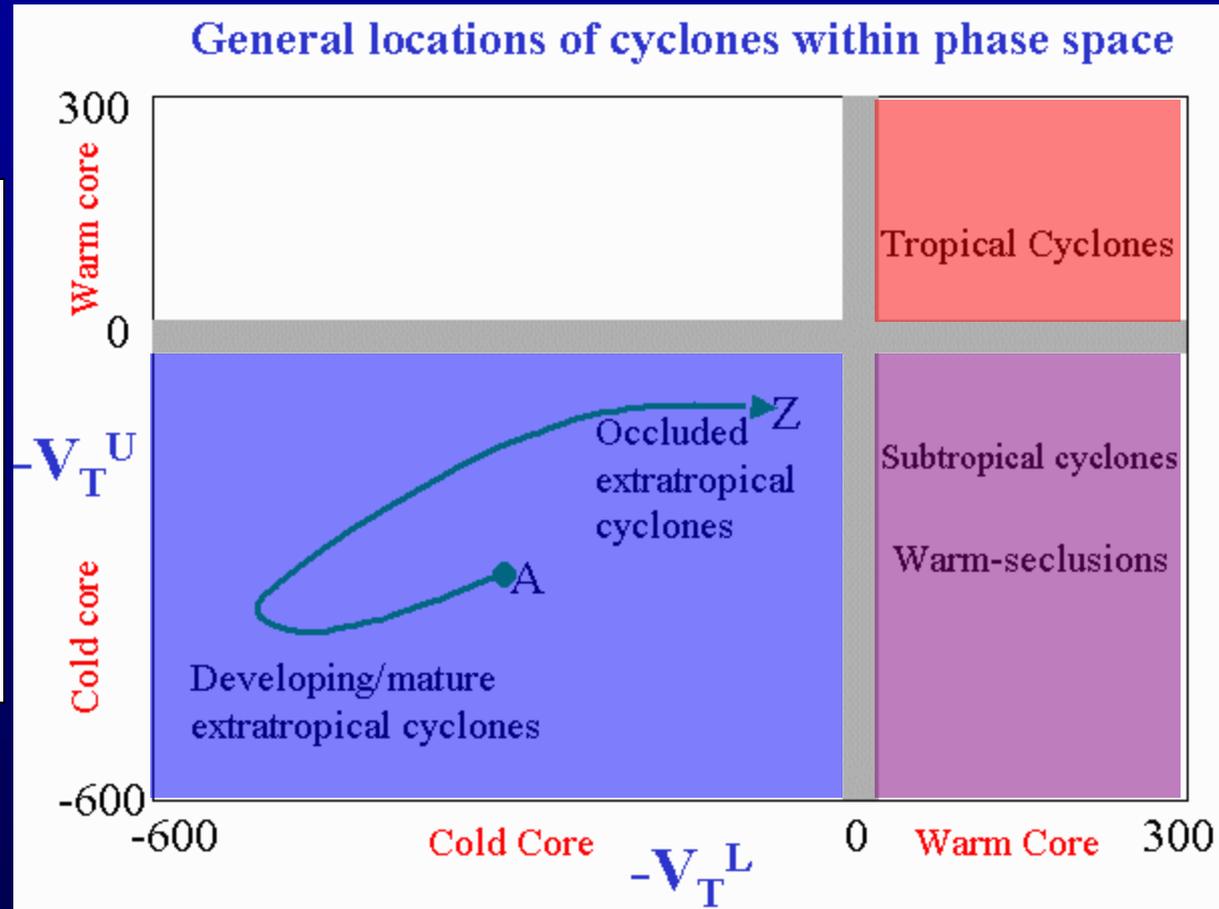
Deep cold core

- $-V_T^L < 0, -V_T^U < 0$
- Extratropical cyclones

Shallow warm core

- $-V_T^L > 0, -V_T^U < 0$
- Subtropical cyclones, warm seclusions

Upper Troposphere



Lower Troposphere



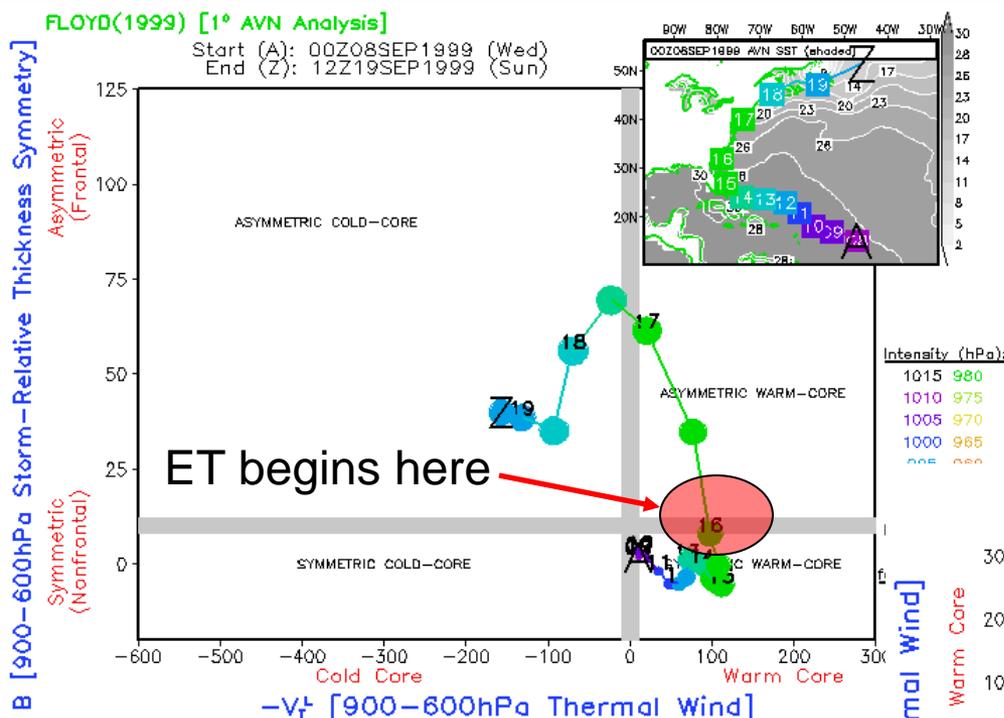
Classification Using Phase Space Diagrams

Evans and Hart (2003)

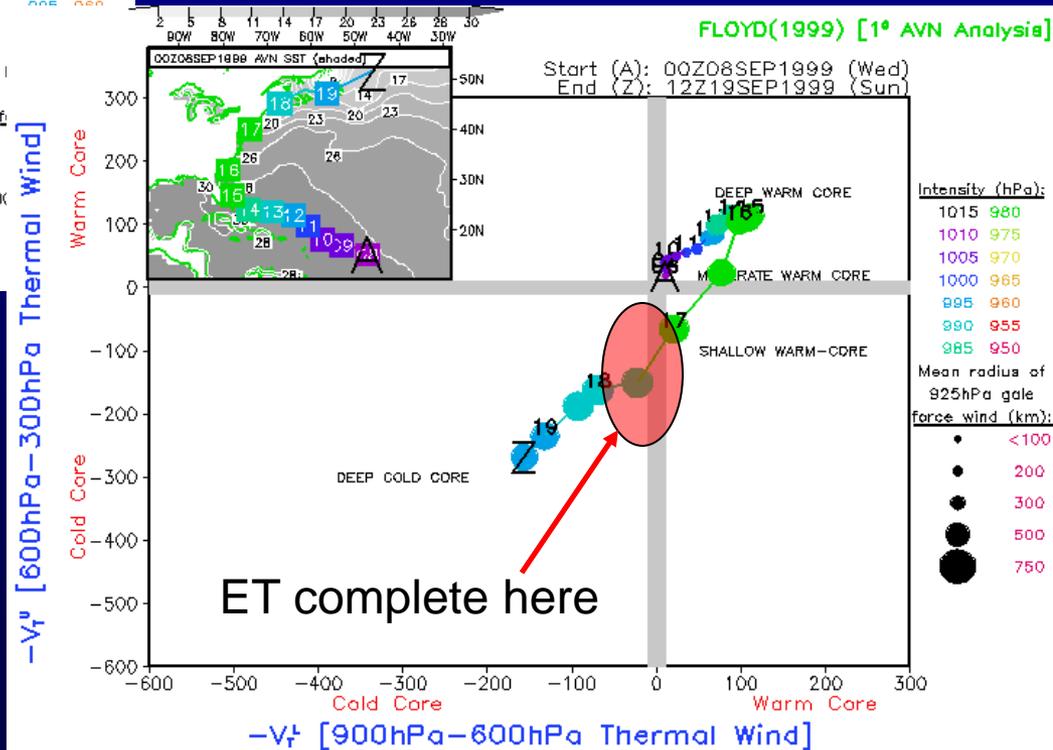


- 61 Atlantic ET events from 1979–1993 examined using ECMWF reanalysis
- Onset of ET: $B > 10$ m
 - Corresponds to onset of low-level frontogenesis
- ET complete: $-V^L_T < 0$
 - Vertical structure is cold-core, i.e., strongest winds aloft
- Mean transition period around 1.5 days

Phase Space Life Cycle Examples



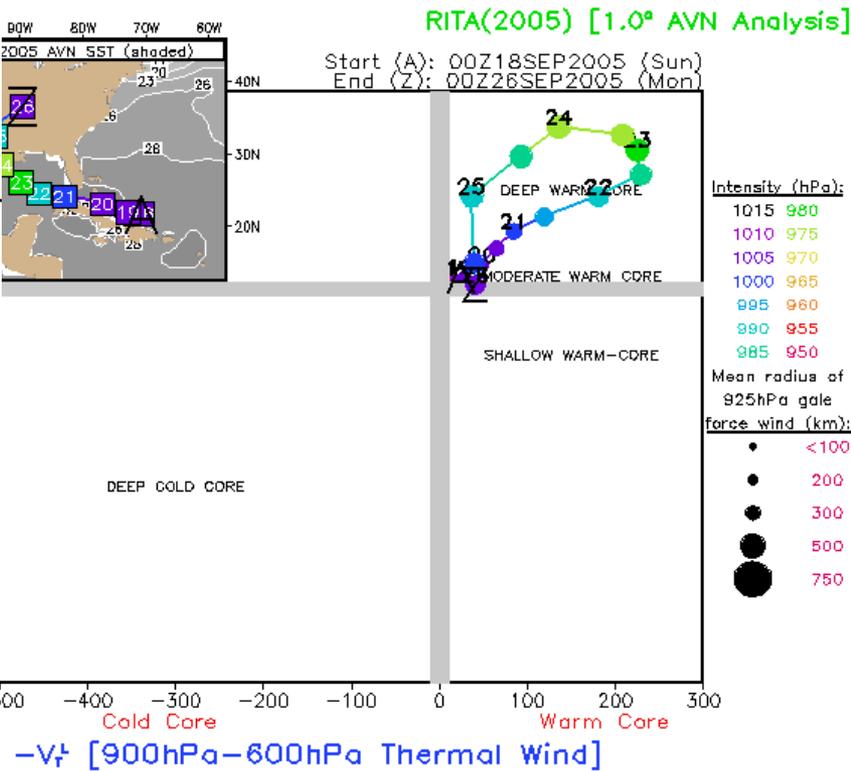
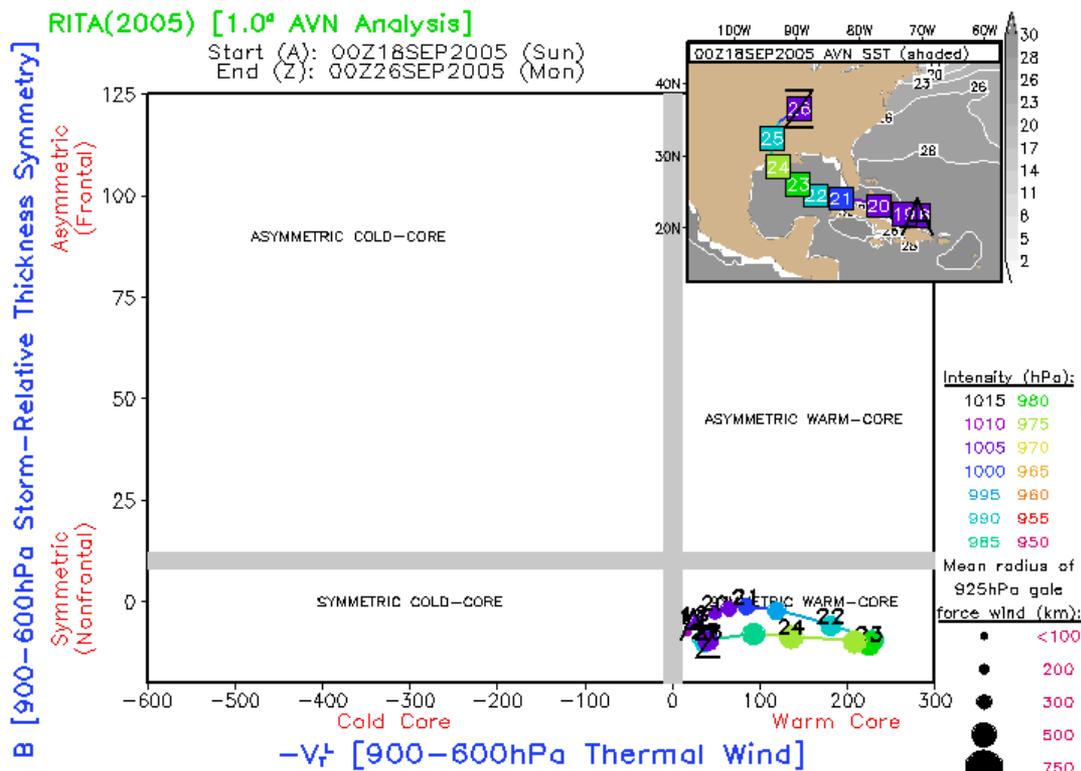
Hurricane Floyd (1999) undergoes
 “typical” ET in phase space
 Initially deep warm core transitions
 to shallow asymmetric warm core
 and then deep asymmetric cold core



Transition of Floyd takes about 24 hours
 Begins early on 16 September
 Ends around 12 UTC 17 September
 (Evans and Hart 2003)

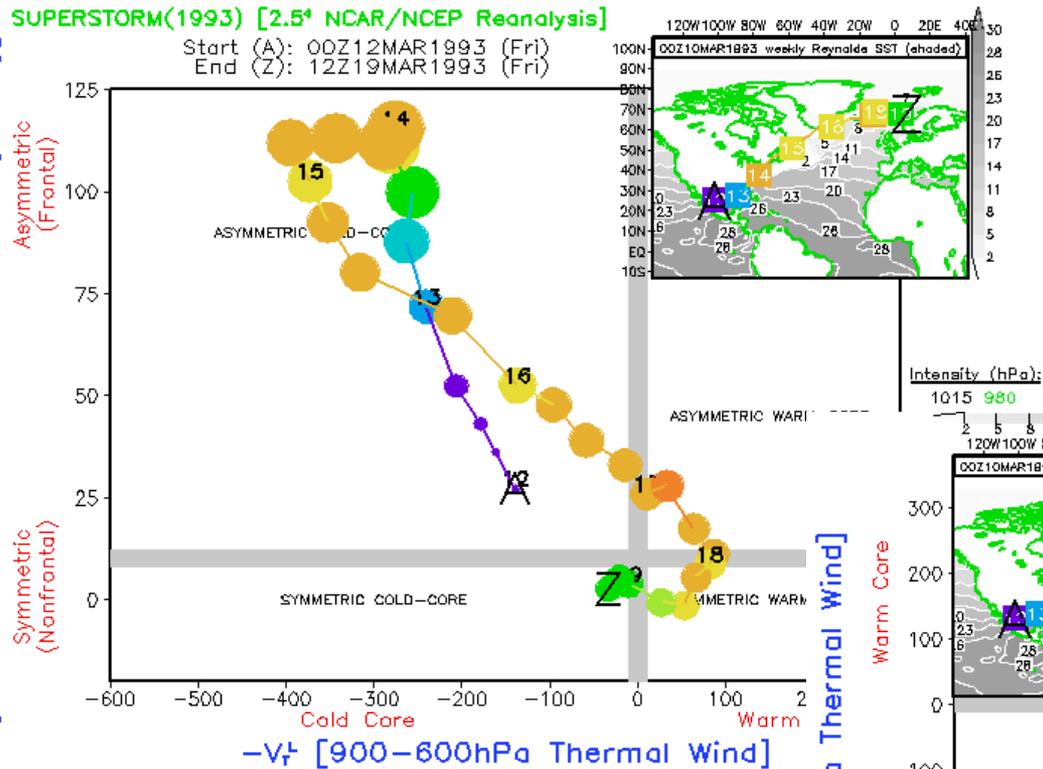
Phase Space Life Cycle Examples

Hurricane Rita (2005)
Major hurricane, developed deep symmetric warm core, did not undergo ET



Phase Space Life Cycle Examples

B [900-600hPa Storm-Relative Thickness Symmetry]

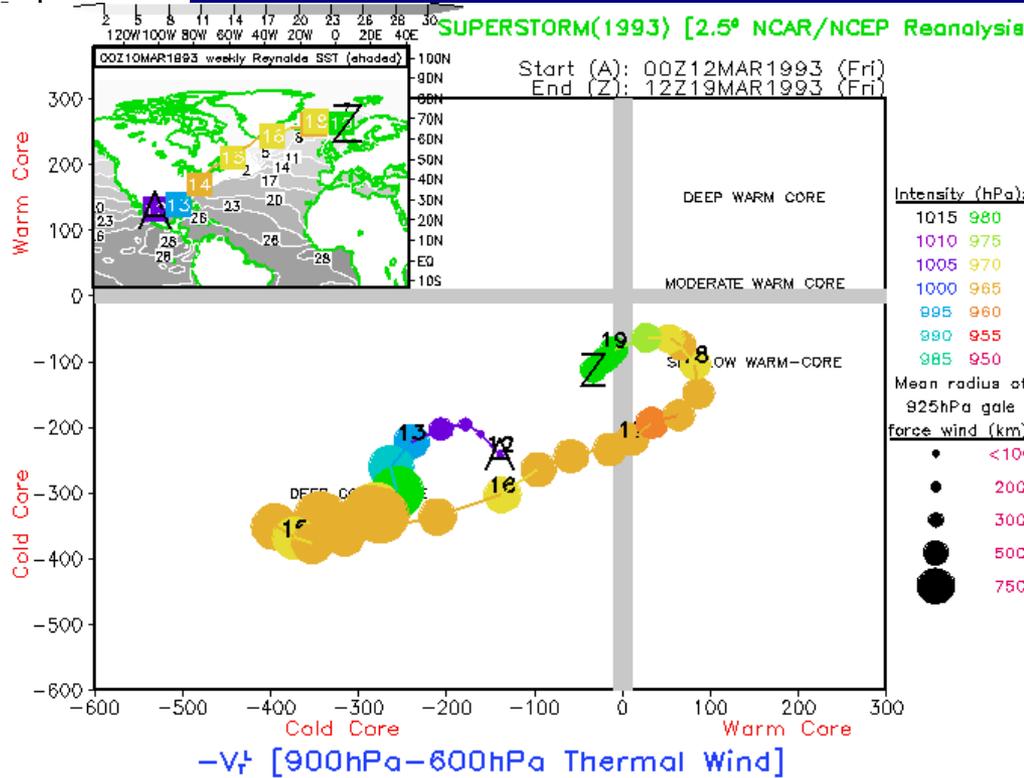


March 1993 Superstorm

Explosive extratropical cyclogenesis, deep and highly asymmetric cold core cyclone

Later acquired warm seclusion characteristics

-V_T [600hPa-300hPa Thermal Wind]



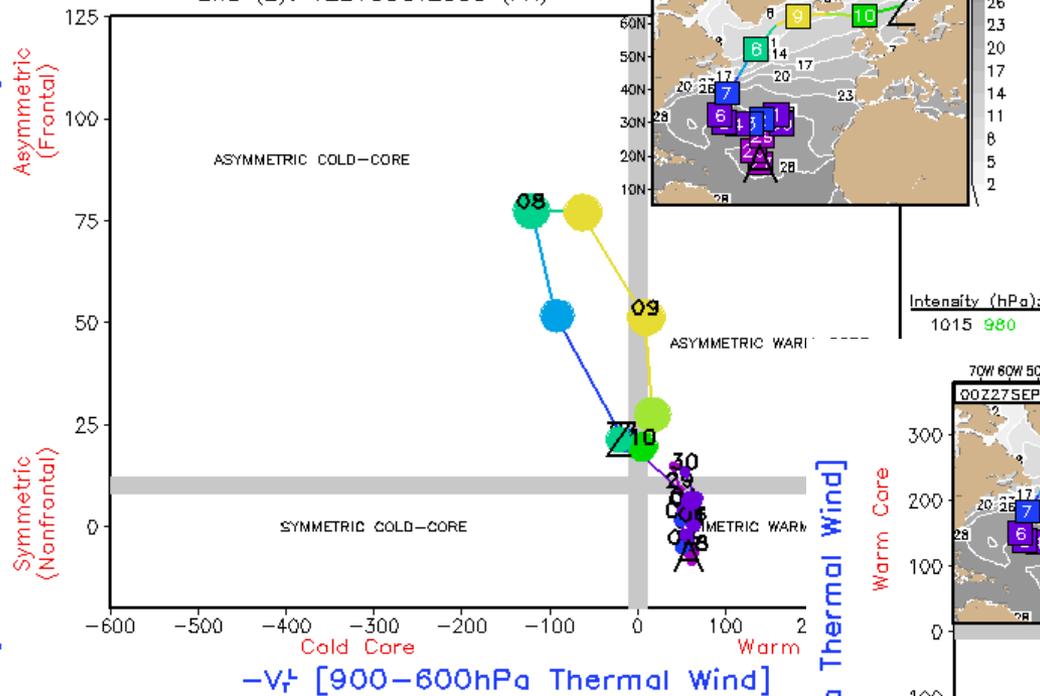


Phase Space Life Cycle Examples



KATE(2003) [1.0° AVN Analysis]

Start (A): 00Z27SEP2003 (Sat)
End (Z): 12Z10OCT2003 (Fri)

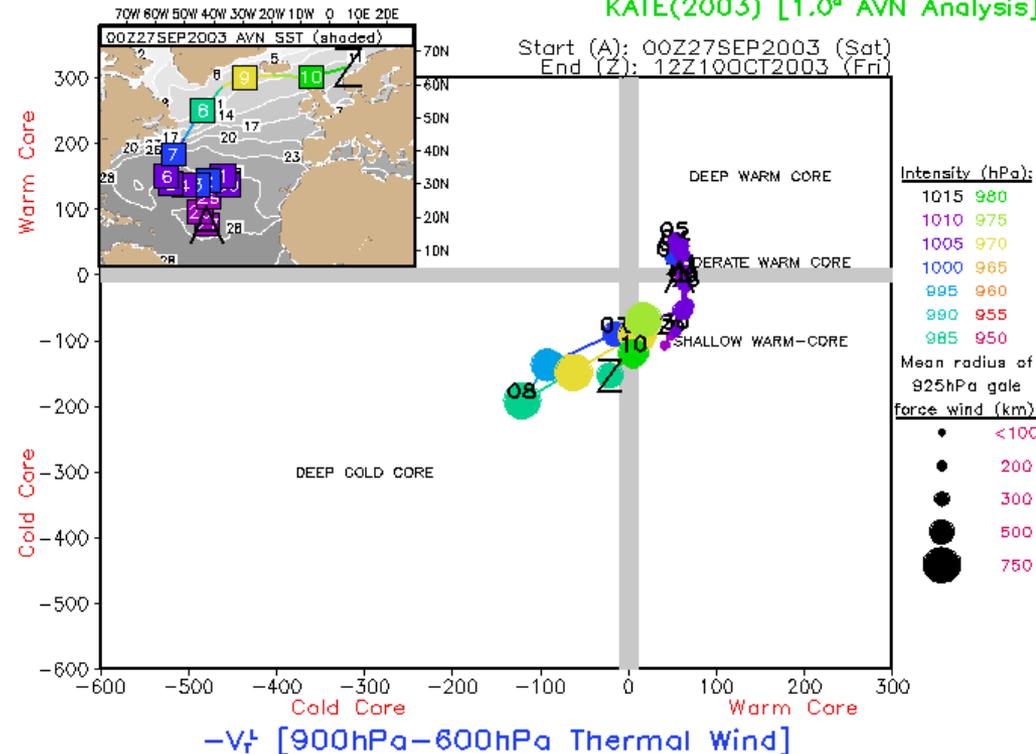


Hurricane Kate (2003): ET interrupted
Begins ET, becoming shallow and less symmetric, only to become more symmetric and with a deeper warm core after moving over warmer water.
Finally undergoes ET later over the north Atlantic

KATE(2003) [1.0° AVN Analysis]

Start (A): 00Z27SEP2003 (Sat)
End (Z): 12Z10OCT2003 (Fri)

-V_T^2 [600hPa-300hPa Thermal Wind]



B [900-600hPa Storm-Relative Thickness Symmetry]

-V_T^2 [900-600hPa Thermal Wind]



Sandy Phase Space Diagram GFS Analysis



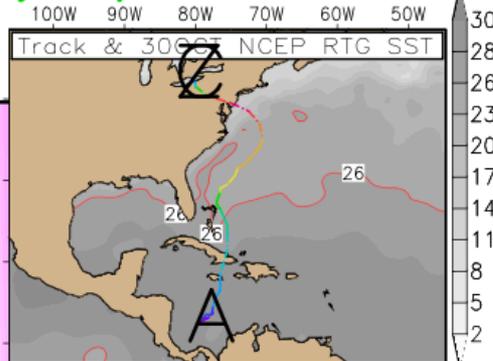
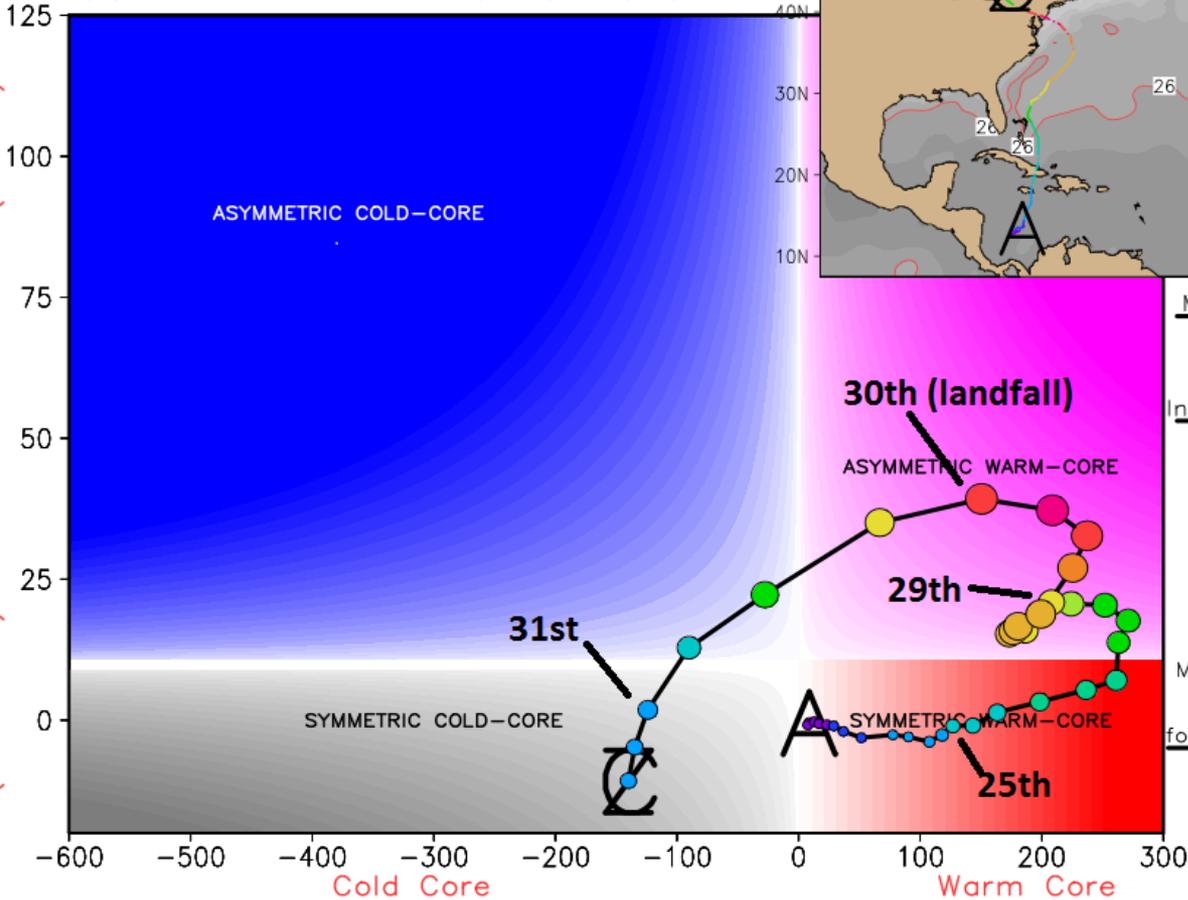
0.5° NCEP GFS (12Z31OCT2012 run) Cyclone #1 (Existing cyclone)

Start (A): 00Z22OCT2012 (Mon) (-228h)
Current (C): 12Z31OCT2012 (Wed) (0h)
End (Z): 12Z31OCT2012 (Wed) (0h)

B [900-600hPa Storm-Relative Thickness Symmetry]

Asymmetric (Frontal)

Symmetric (Nonfrontal)



Marker Style:

- Analysis: ●
- Forecast: ⊗

Intensity (hPa):

1015	980
1010	970
1000	960
990	950

Mean radius of 925hPa gale force wind (km):

- <100
- 200
- 300
- 500
- 750

NOTE: A 24hr running mean smoother is applied to the CPS trajectory.



**But, it isn't always that
simple...**



Phase Space Caveats

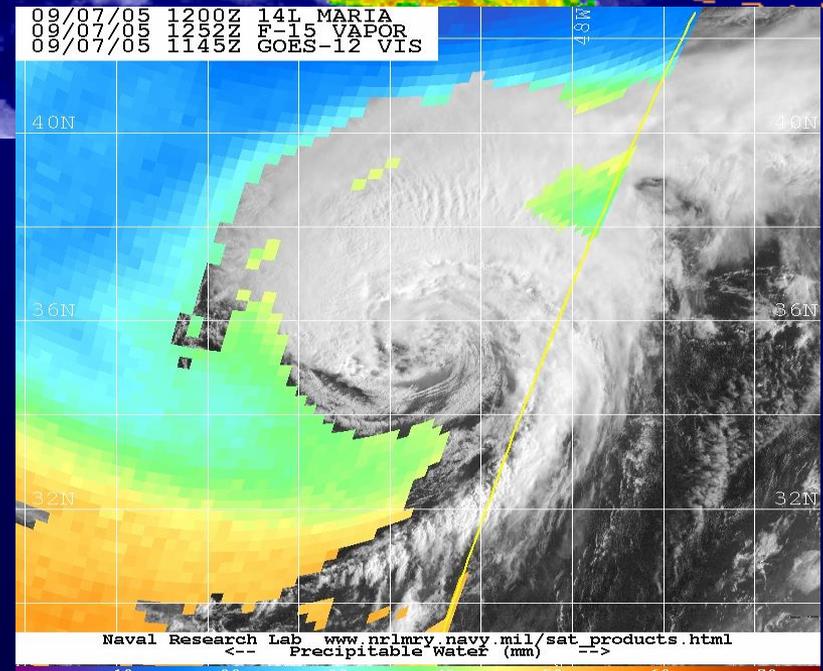
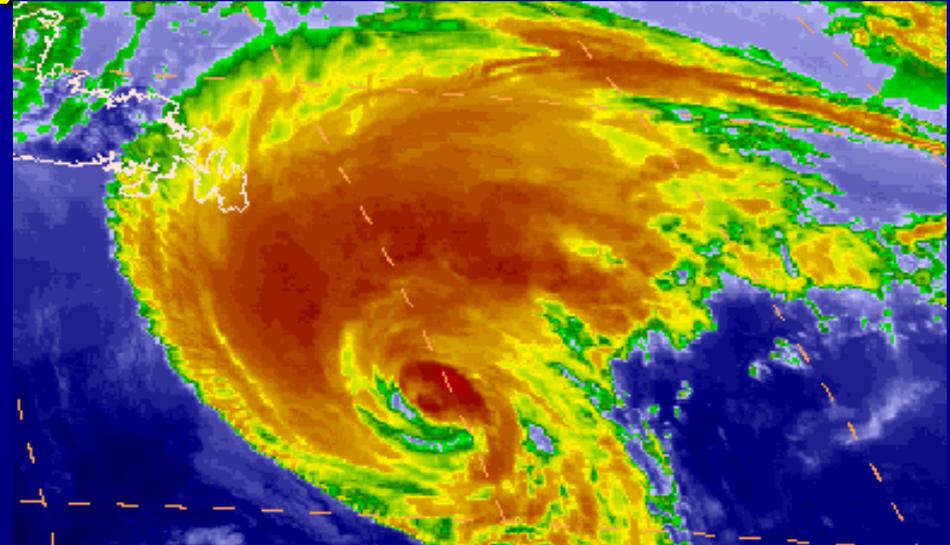
- Quality of phase space diagrams depends on the model analyses and forecasts used to create them
 - Errors in model analysis of cyclone intensity and structure and forecasts of track, intensity, and structure will manifest themselves in the phase space
- Model fields used to create diagrams still relatively coarse (0.5° to 1° for global models), so small scale structure not captured
- Impacts of TC bogusing
 - Stronger warm core noted for “active” TCs (i.e., advisories being written)
 - Better representation of purely tropical cyclone before ET has started
 - Warm core persists once ET has begun – may never show transition to cold-core cyclone



Analysis and Short Term Forecast (0–6 h) Tools



- Where is this cyclone in the ET process?
 - Satellite/radar imagery
 - Intensity estimates
 - AMSU, Dvorak, Hebert-Poteat satellite classifications
 - Observations of surface wind field (ship/buoy, land, satellite)
 - Surface and upper-air analysis
 - Model analyses
 - SHIPS classification





Forecast Tools (Beyond 6 h)

- NWP model guidance
 - Is analyzed structure of cyclone undergoing ET consistent with what you currently see in observations? If not, be wary!
 - Are critical features (upper-level troughs/jets, fronts, extratropical cyclones) properly represented in the model analysis?
 - What signals of ET do you see in model forecasts? Expansion of wind field, increased thermal advections, intensification through baroclinic processes?
 - How much spread is there between the various models? Do all models show ET occurring, only some? Timing different?
 - Cross sections of potential vorticity, vorticity, moisture, wind, to examine vertical structure of cyclone
 - SHIPS model classification
- Conceptual Models of ET and climatology
 - How does what you see in observations and model guidance fit into conceptual models of ET?
 - Is this event an outlier in terms of the ET climatology for the basin?



SHIPS Model Phase Information



- Classification scheme uses the following variables:
 1. SST
 2. Storm translational speed
 3. Cold pixel counts from IR imagery (less important during forecast period)
 4. Difference between 300- and 500-hPa tangential winds within 500 km of storm center (related to upper-level warm/cold core)
 5. 150-hPa temperature (related to tropopause height)
 6. 850–700-hPa temperature gradient within 500 km of storm center
 7. 850–200-hPa vertical shear
- Trained on NHC best track 1982-2010
 - Correct classification 91% of time at $t=0$
 - Most reliable for distinguishing between Tropical and Extratropical
 - Less reliable for Subtropical



SHIPS Model Phase Information



* ATLANTIC SHIPS INTENSITY FORECAST *
 * GOES AVAILABLE, OHC AVAILABLE *
 * BERTHA AL032014 08/04/14 12 UTC *

TIME (HR)	0	6	12	18	24	36	48	60	72	84	96	108	120
V (KT) NO LAND	70	74	77	78	78	79	78	77	71	68	60	50	37
V (KT) LAND	70	74	77	78	78	79	78	77	71	68	60	50	37
V (KT) LGE mod	70	76	80	81	80	76	74	67	57	47	41	39	39
Storm Type	TROP	EXTP	EXTP	EXTP	EXTP	EXTP							
SHEAR (KT)	21	18	20	22	22	33	45	55	54	62	60	65	71
SHEAR ADJ (KT)	1	0	0	-1	-4	0	5	-2	2	-3	0	-7	-1
SHEAR DIR	337	322	305	299	280	254	237	230	236	238	244	256	263
SST (C)	29.0	28.8	28.4	27.9	27.4	27.5	26.6	24.4	19.2	15.0	14.4	15.9	14.1
POT. INT. (KT)	153	151	145	138	132	134	124	105	81	71	70	72	70
ADJ. POT. INT.	141	138	132	124	117	119	111	93	74	68	67	68	67
200 MB T (C)	-54.1	-53.6	-53.3	-53.4	-53.6	-53.6	-53.9	-53.7	-54.2	-54.8	-54.1	-53.9	-53.1
TH_E DEV (C)	10	10	9	8	7	7	5	4	2	0	0	0	0
700-500 MB RH	55	59	62	62	63	55	55	58	62	60	60	63	71
GFS VTEX (KT)	11	11	12	11	11	14	16	19	21	27	29	29	26
850 MB ENV VOR	-77	-40	-35	-41	-29	-9	51	98	112	101	94	127	170
200 MB DIV	62	80	82	28	36	69	116	111	76	94	53	41	17
700-850 TADV	21	23	18	18	24	10	18	19	24	29	-4	-37	-27
LAND (KM)	644	672	612	503	437	593	528	409	256	256	590	961	1335
LAT (DEG N)	26.8	28.4	30.0	31.7	33.3	36.1	38.8	41.7	44.5	46.6	47.8	48.8	49.7
LONG (DEG W)	73.6	73.5	73.3	72.6	71.9	69.0	64.5	59.6	54.7	49.7	44.9	39.9	34.7
STM SPEED (KT)	15	16	17	18	17	20	23	23	22	19	17	18	17
HEAT CONTENT	35	33	17	12	7	23	13	0	0	0	0	0	0



Summary (1 of 2)



- ET is a complex process, involving interactions between phenomena on multiple scales
- ET represents an operational challenge because cyclones exist on a continuum – they don't fit into discrete bins!
- Operational decisions during ET often take more than cyclone structure into consideration
 - Impacts, watch/warning complications, coordination
- Structural changes during ET have large impacts on sensible weather
 - Redistribution of clouds and heavy precipitation
 - Expansion of wind field has large impact for land and marine interests
 - Dramatic increase in forward speed



Summary (2 of 2)



- Conceptual models largely based on satellite imagery and large-scale model analyses
- Cyclone phase space diagrams show evolution of structure based on model analyses and forecasts
 - Neither may be able to resolve small scale features important to individual ET cases
- Most composites and statistics based on relatively short data records and small sample sizes, so beware of outliers!
- Examine NWP model guidance carefully
 - Differences in initial structure and evolution of both the TC and larger-scale flow will can result in widely varying forecasts of cyclone structure between the models, and from run to run
 - Does initial analysis capture the correct structure of the TC and other features important during the ET process?



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- Hart, R. E., 2003: A cyclone phase space derived from thermal wind and thermal asymmetry. *Mon. Wea. Rev.*, **131**, 583–616.
- , and J. L. Evans, 2001: A climatology of extratropical transition of Atlantic tropical cyclones. *J. Climate*, **14**, 546–564.
- , J. L. Evans, and C. Evans, 2006: Synoptic composites of the extratropical transition life cyclone of North Atlantic tropical cyclones: Factors determining posttransition evolution. *Mon. Wea. Rev.*, **134**, 553–578.
- Jones, S. C., and coauthors, 2003: The extratropical transition of tropical cyclones: Forecast challenges, current understanding, and future directions. *Wea. Forecasting*, **18**, 1052–1092.
- Klein, P. M., P. A. Harr., and R. L. Elsberry, 2000: Extratropical transition of western North Pacific tropical cyclones: An overview and conceptual model of the transformation stage. *Wea. Forecasting*, **15**, 373–396.
- Wood, K. M., and E. A. Ritchie, 2014: A 40-year climatology of extratropical transition in the eastern North Pacific. *J. Climate*, **15**, 5999–6015.



Forecast Exercise





Exercise



- Break into three groups by row
- The situation
 - 00 UTC 22nd of the month
 - Satellite imagery and current status of cyclone
- Each group will have the following data from a different global model
 - Analyzed phase space diagram through 00 UTC 22nd
 - Model forecast fields every 12 h through 96 hours (00 UTC 26th)
- Using only the data available to them each group will estimate
 - Beginning time of ET
 - When the system would be declared extratropical
- Group 1: GFS
- Group 2: NOGAPS
- Group 3: UKMET



Exercise



- Next, each group will be given the analyzed cyclone phase space diagrams from the model analysis for the entire lifecycle of the cyclone and satellite imagery
- Compare your forecasts to what the model analysis and imagery shows



Exercise



- Finally, we'll look at the satellite imagery and operational NHC classification for this system and compare